

Geochemistry and Mantle Source(s) for Carbonatitic and Potassic Lavas from SW Uganda

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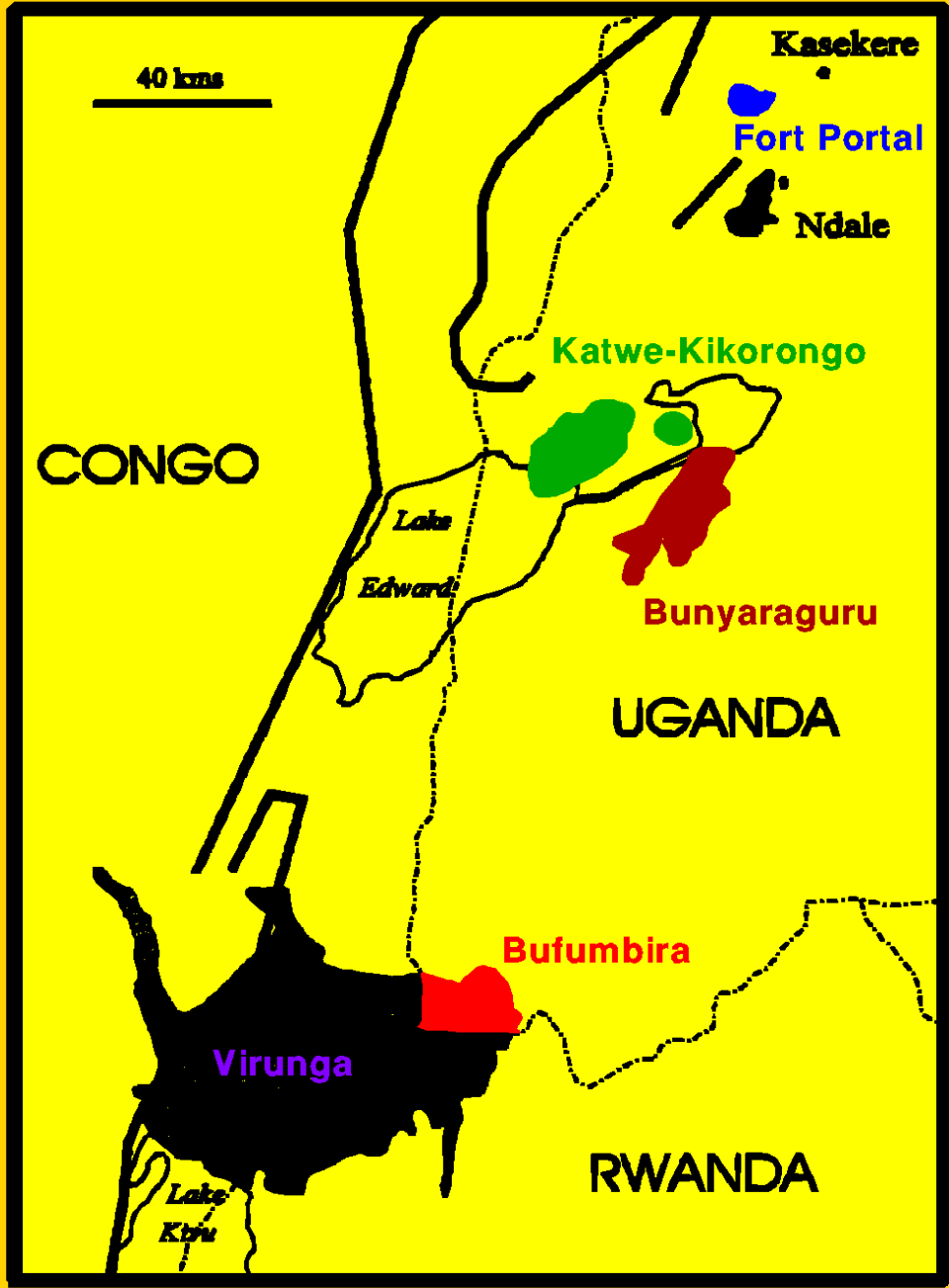
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Bunyaraguru
Olivine-bearing
tephras & rare
lavas. Leucite +
augite (ugandite),
augite + kalsilite
(mafurite) and
melilite + leucite
(katungite)

Bufumbira
Basanite, leucitite,
leucite-phonolite,
latite & trachyte



Fort Portal
Extrusive
carbonatites

**Katwe-
Kikorongo**
Olivine-melilitite
and feldspathoidal
cpx-rich tephras &
subordinate flows

Fort Portal

Tuff cone and
crater lake



Tuff cone

Fort Portal

Quarry -
“flaggy” tuff

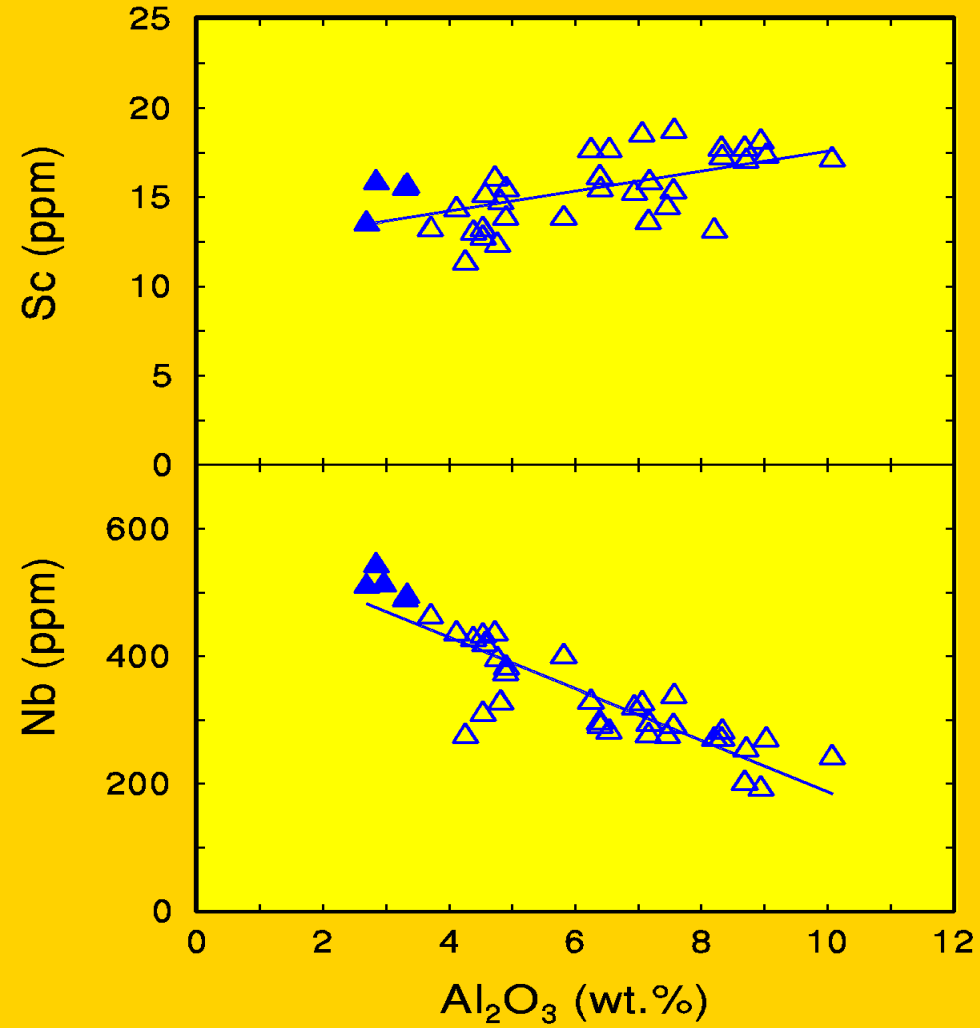
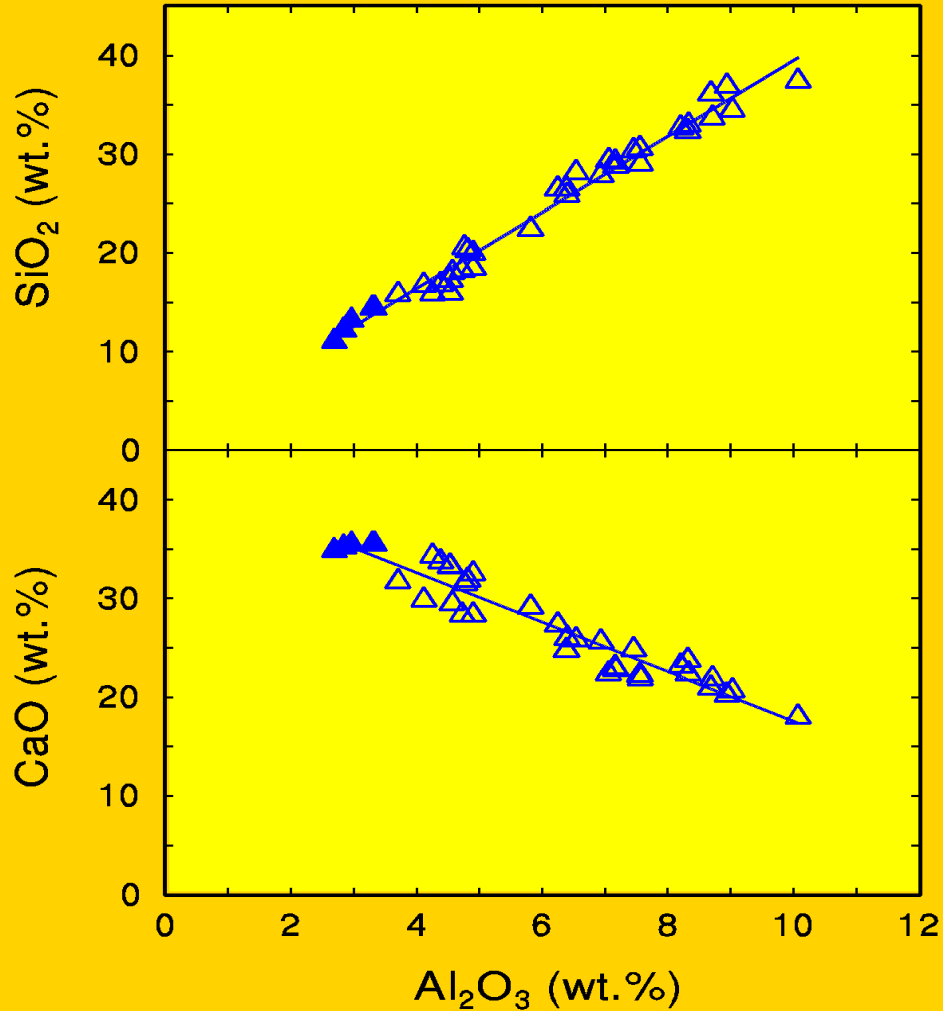


Lapilli-tuff

Ash-tuff

Mixing lines for Fort Portal tuffs and lavas

Extrapolation to $\text{Al}_2\text{O}_3 = 0\%$ yields an estimate of the original composition of the carbonatite magma.



- Crustal xenoliths are ubiquitous in the Fort Portal tuffs and lavas.
- The Fort Portal tuffs and lavas can be regarded as mixtures of carbonatite magma and Precambrian crustal xenoliths.
- Plots of various elements versus Al (or Si) yield straight-lines which represent variable mixtures of crustal xenoliths and carbonatite magma.
- The original composition of the carbonatite magma can be determined by extrapolation to zero wt.% alumina.

Katwe-Kikorongo

Crater

Rim is composed of tuffs and
agglomerates



Crater lake



Katwe-Kikorongo



Subaqueous tuffs

Bufumbira



Trachyte plug



Mgahinga and Sabinio
volcanoes

Bufumbira

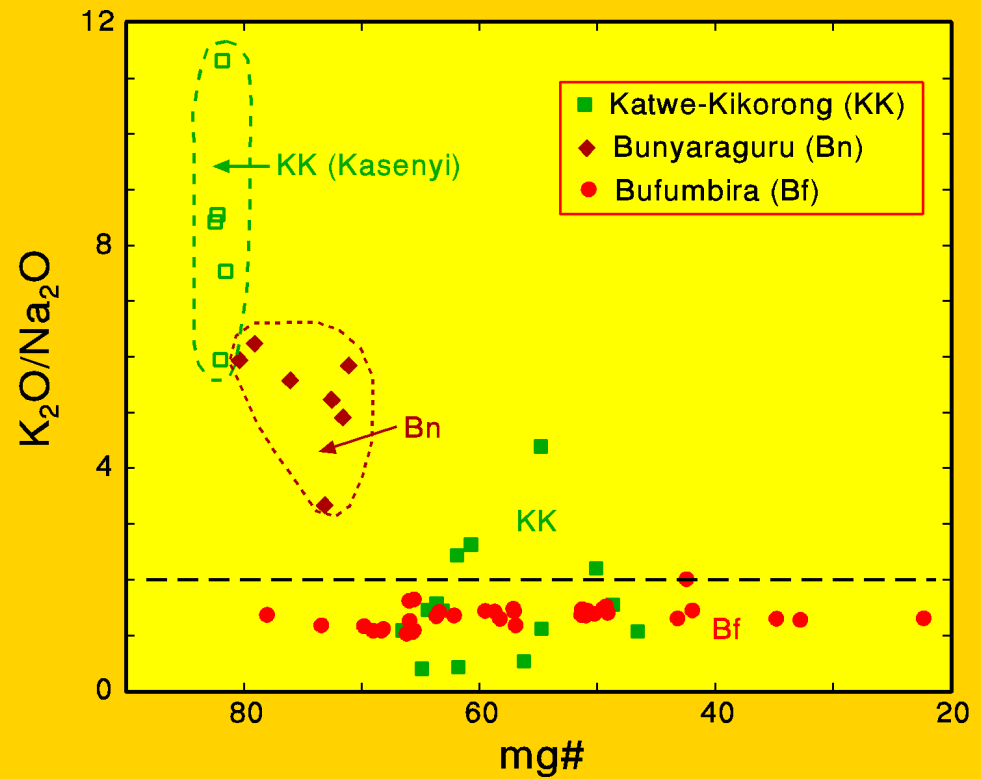
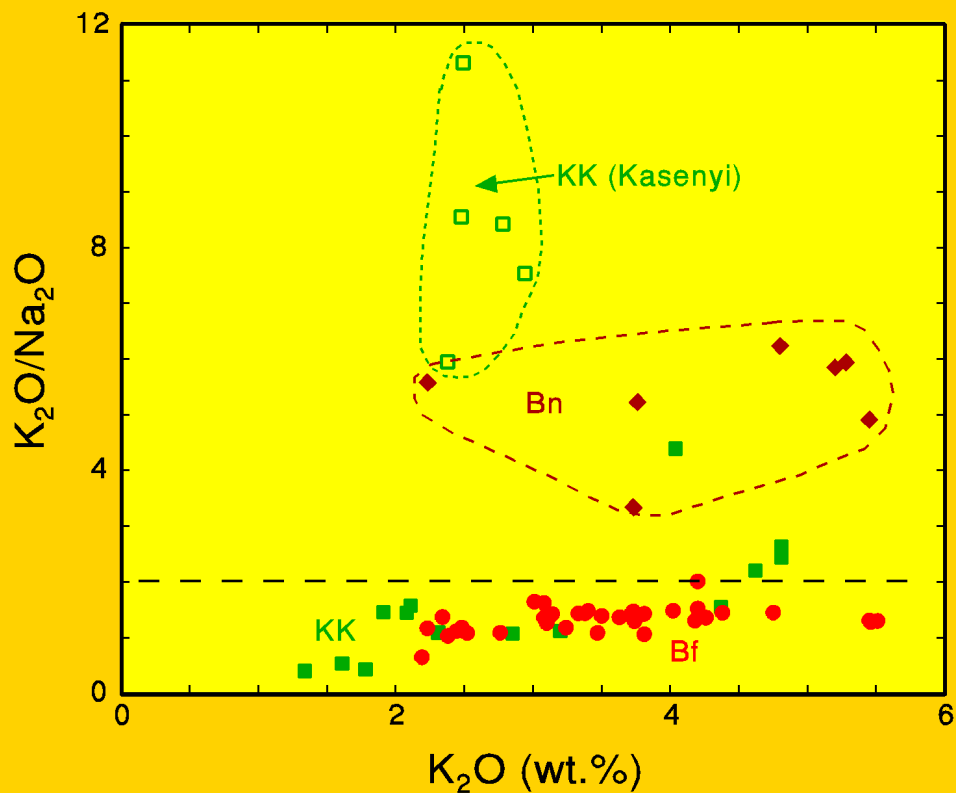
Leucite basalt

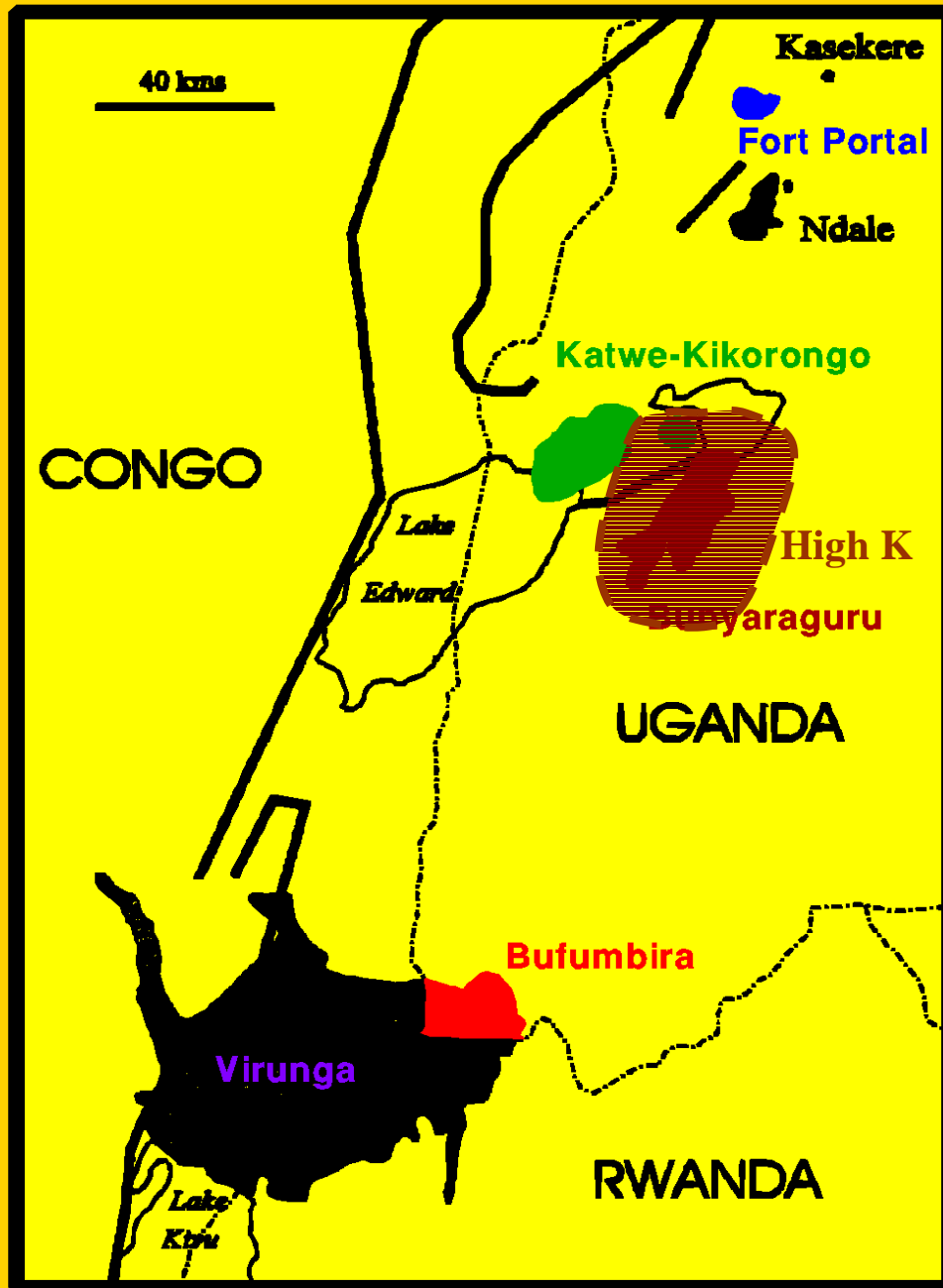


Lapilli tephra

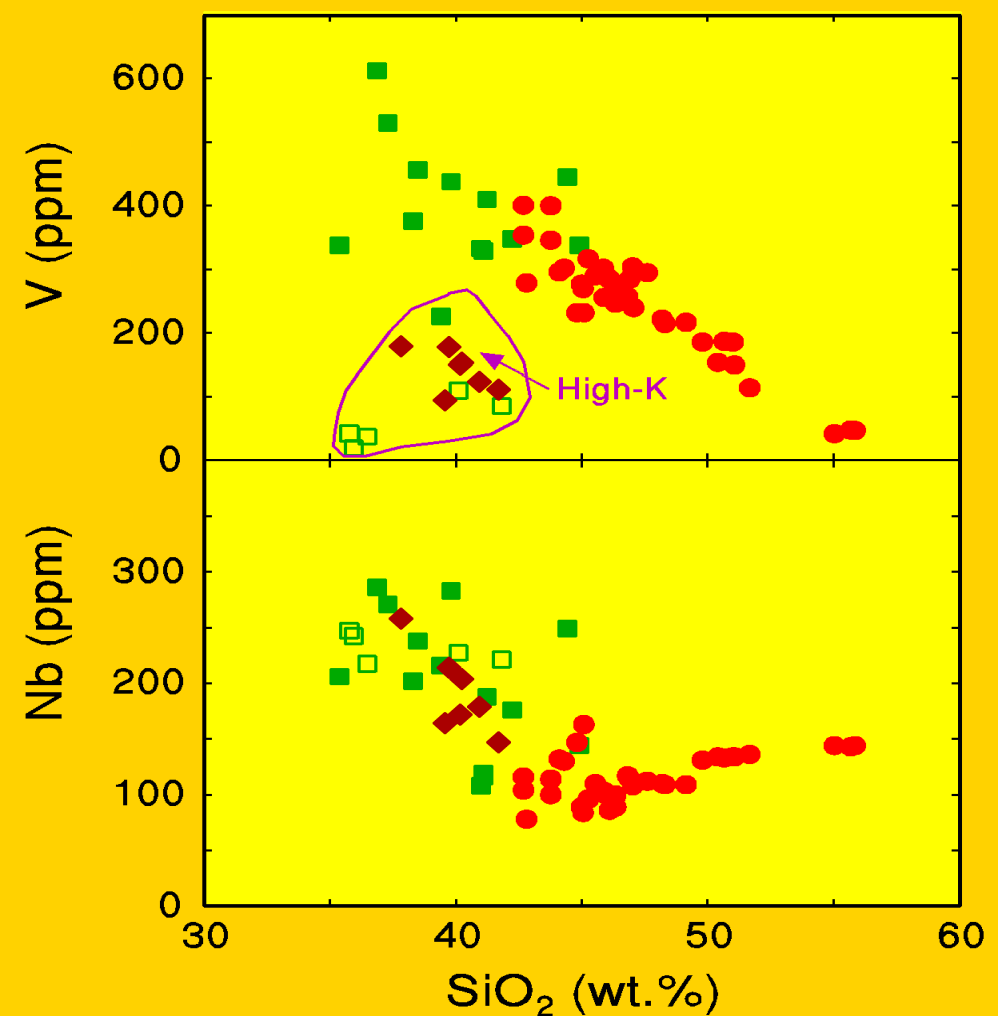
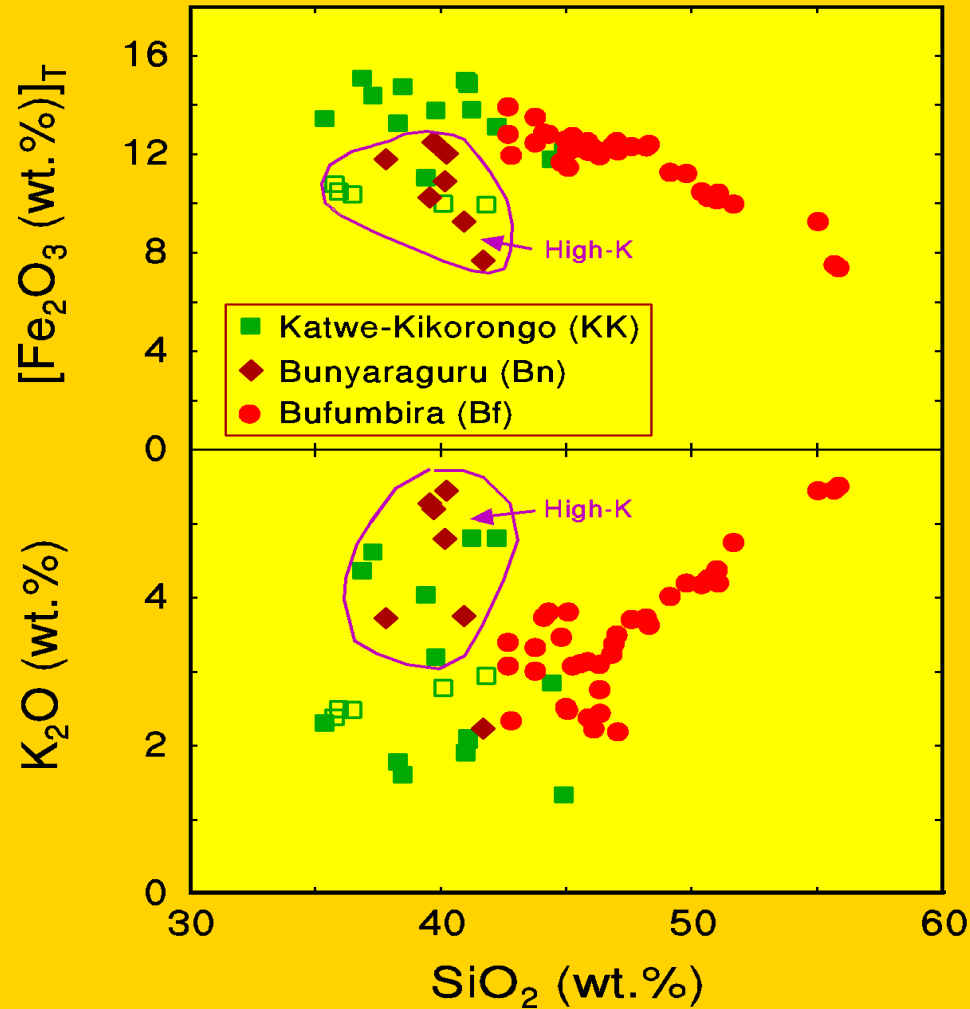


K_2O - Na_2O relationships for Katwe-Kikorongo, Bunyaraguru and Bufumbira. Note the strongly potassic character of the Bunyaraguru and Katwe-Kikorongo (Kasenyi) samples.

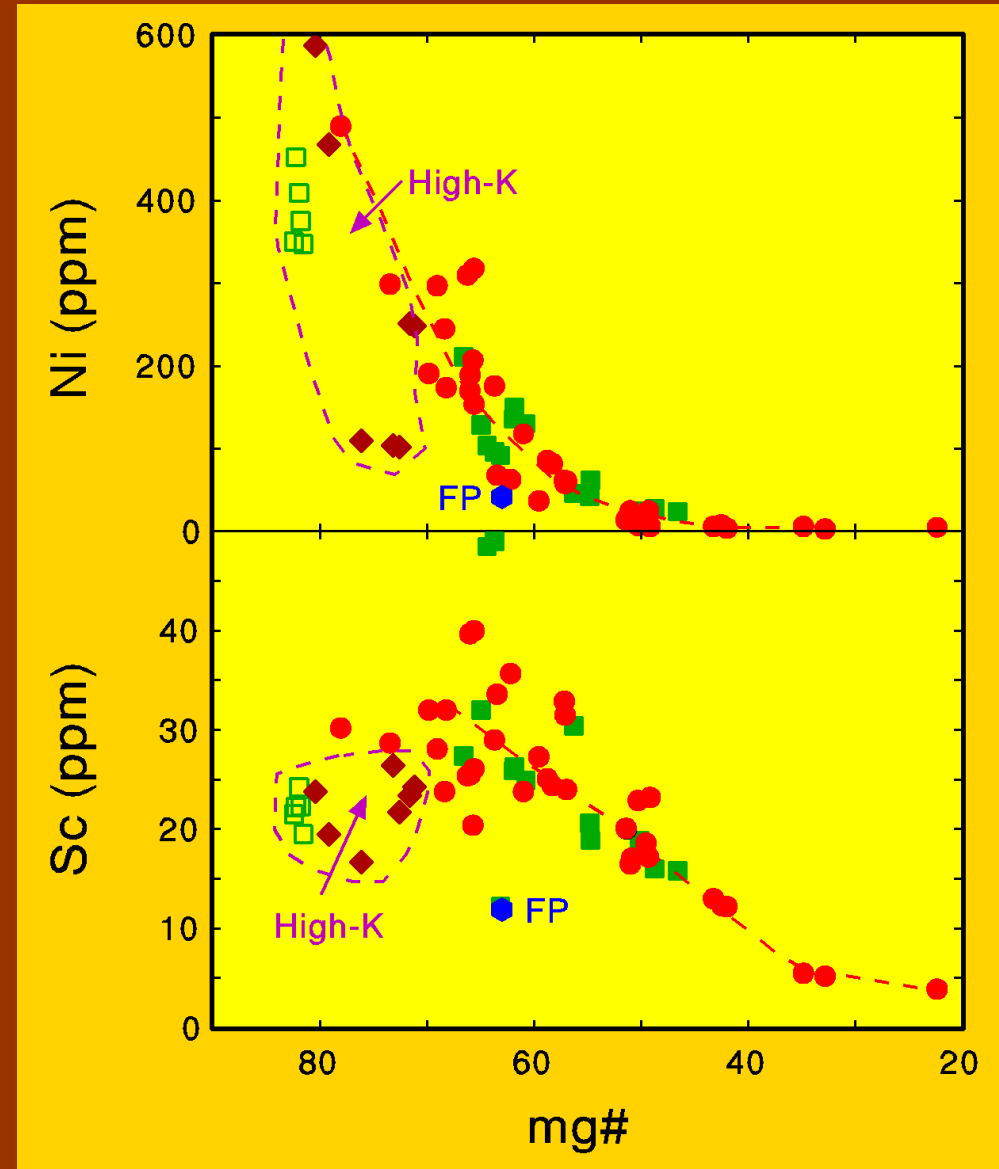
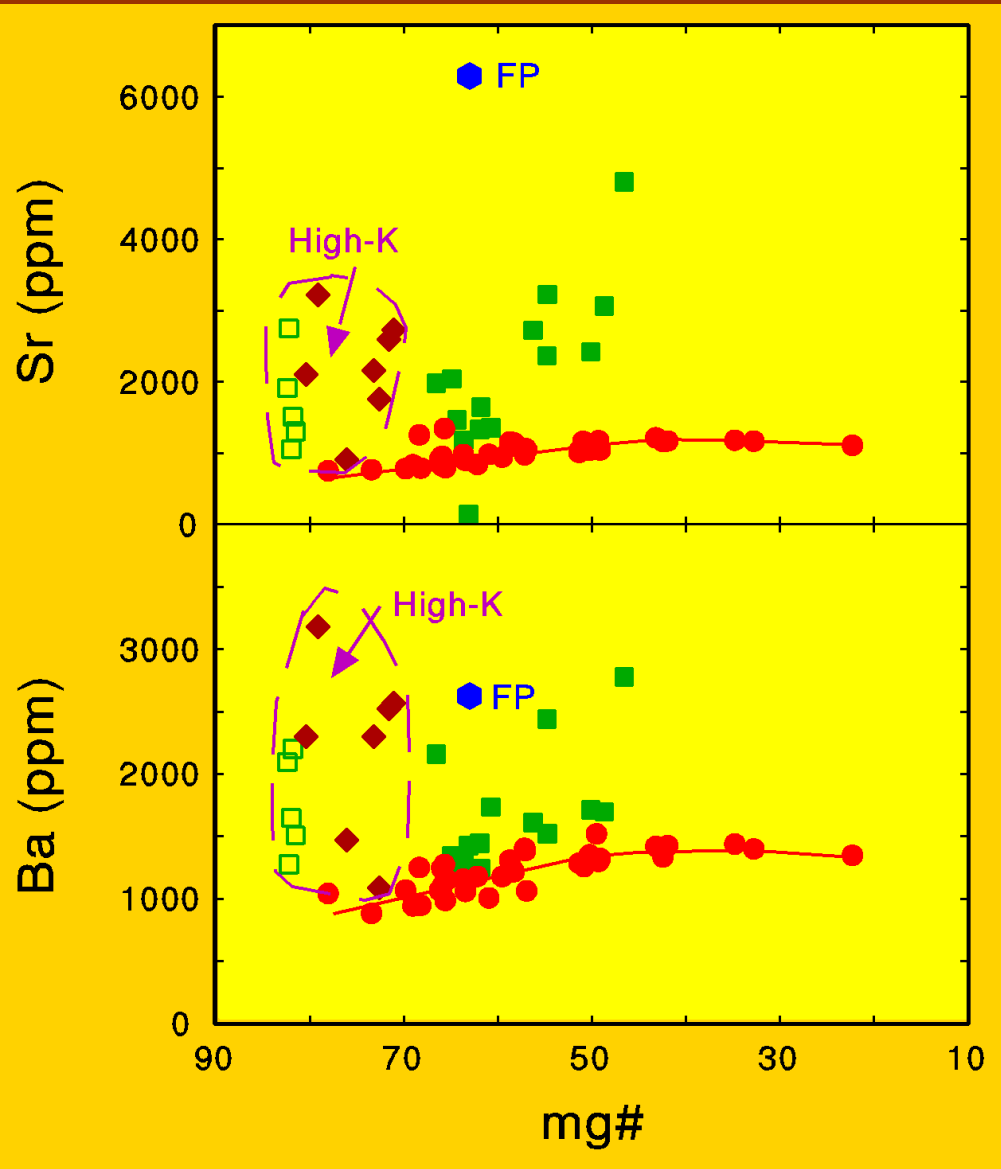




Selected major and trace elements versus SiO_2 . Note that the “potassic” samples plot in distinctly different areas from the rest of Katwe-Kikorongo and Bufumbira samples. Declining V with increasing SiO_2 suggests magnetite fractionation.

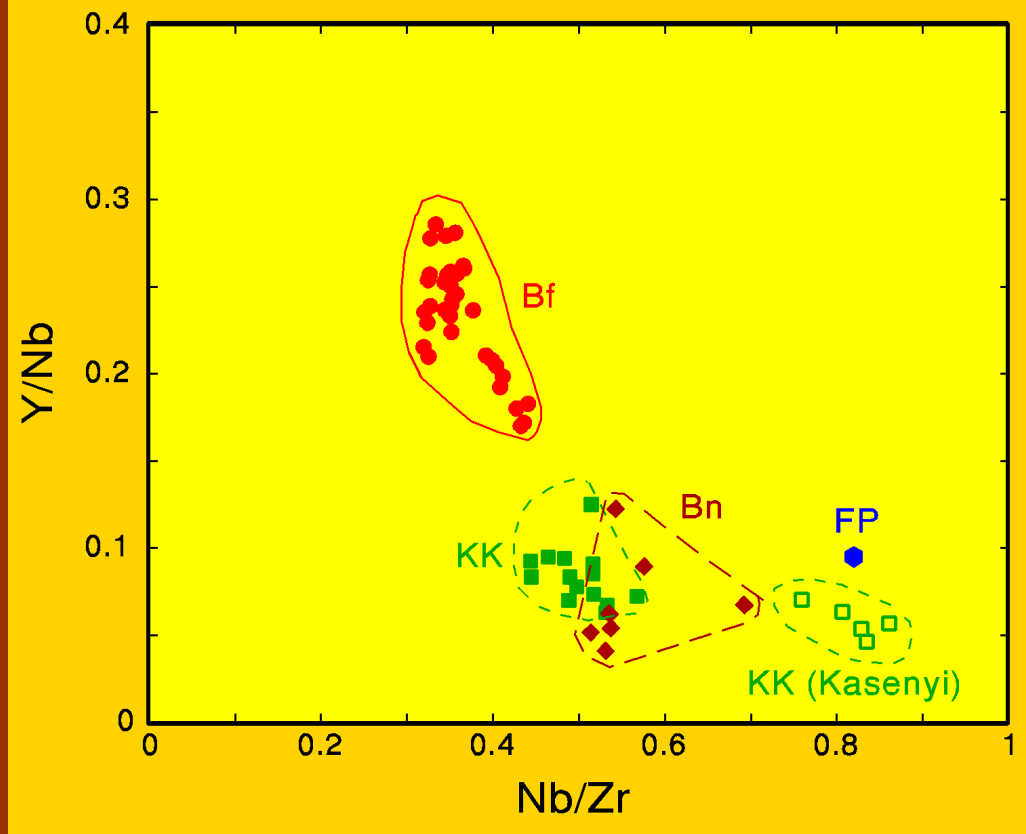
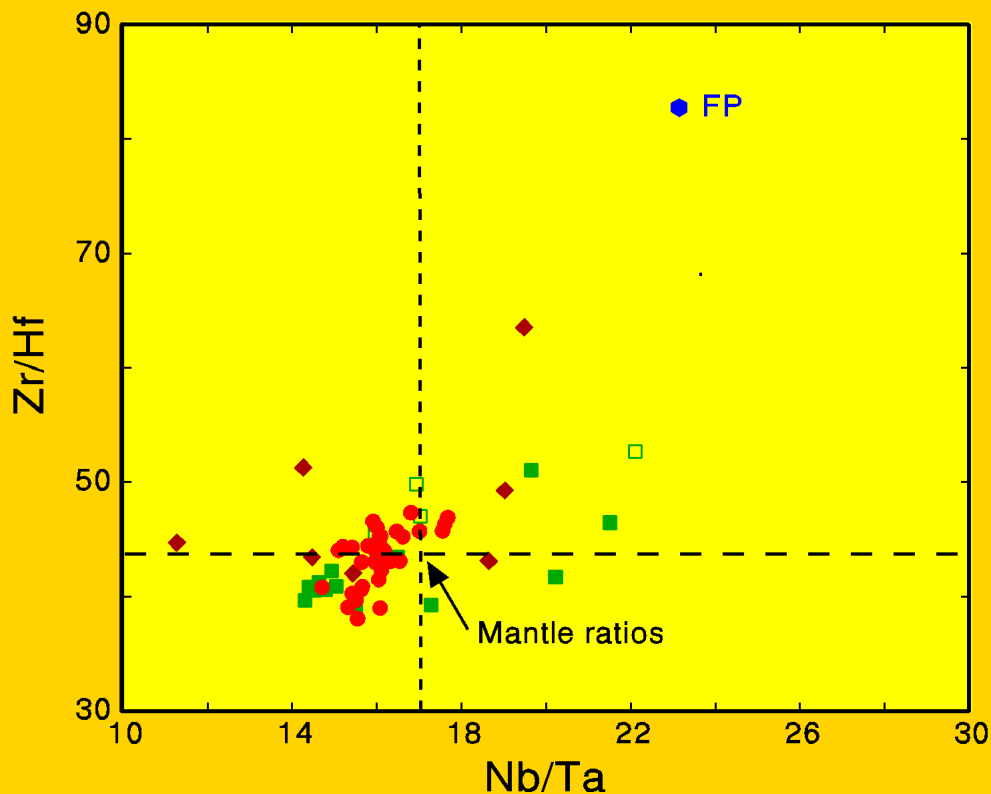


Selected trace element vs mg#. Note that the high-K samples are significantly enriched in Sr and Ba and relatively depleted in Ni and Sc.



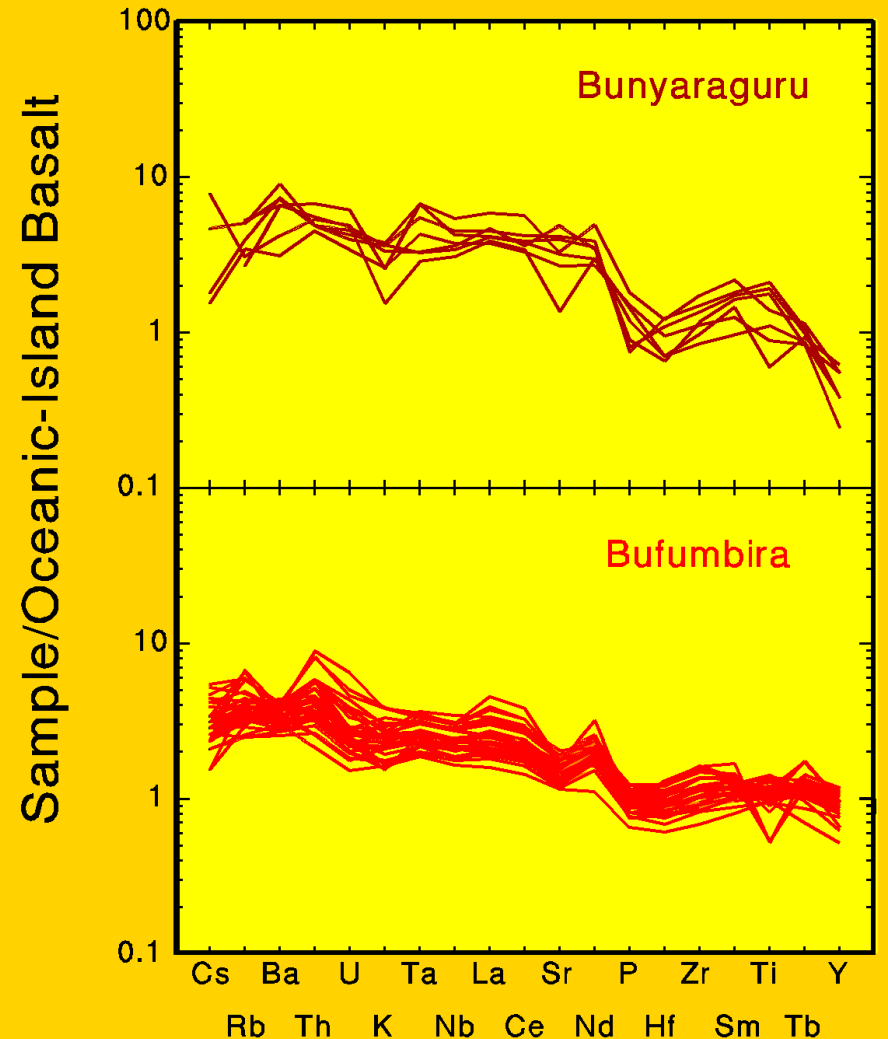
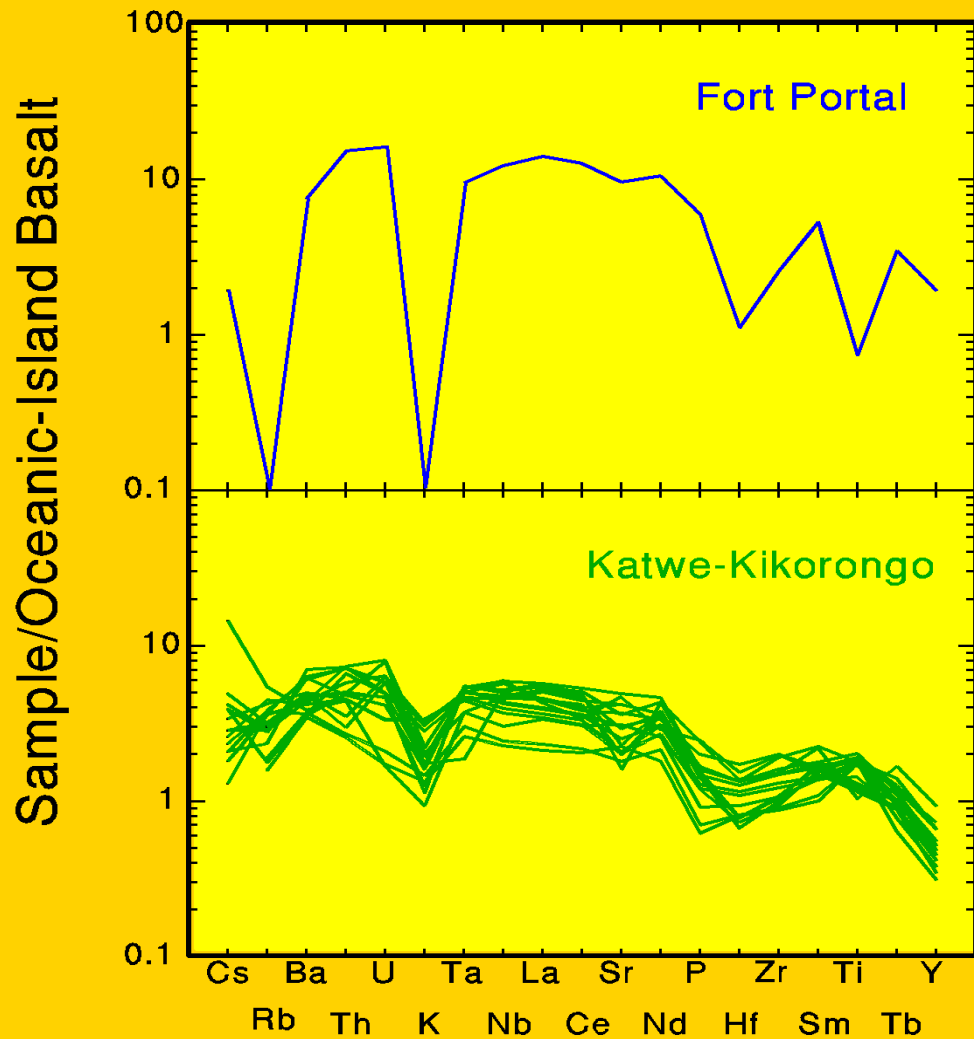
- The high-K volcanics show chemical signatures significantly different from the other volcanics, e.g., higher K_2O , Ba and Sr and lower V, Ni and Sc at similar SiO_2 content and mg#.
- A regular increase in Sr and Ba content, and then a constant concentration, with decreasing mg# indicates that the feldspar minerals were not important fractionating phases.
- The regular decrease in Ni (olivine) and Sc (clinopyroxene) with increasing mg# indicates that these minerals were important fractionating phases.

Each volcanic area plots in a distinct region on a Y/Nb vs Zr/Hf diagram suggesting different mantle source regions.

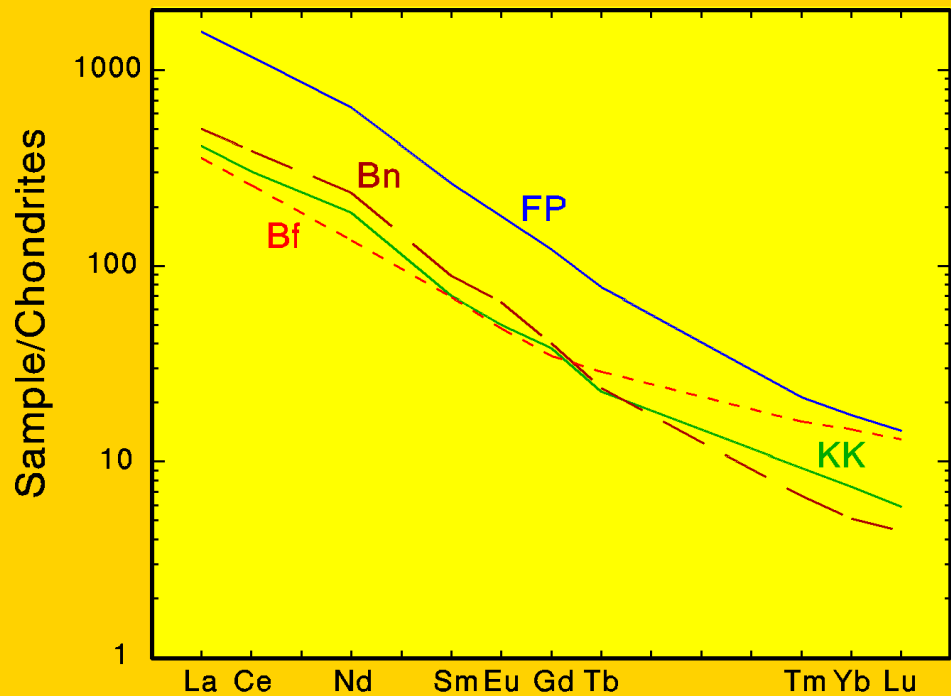
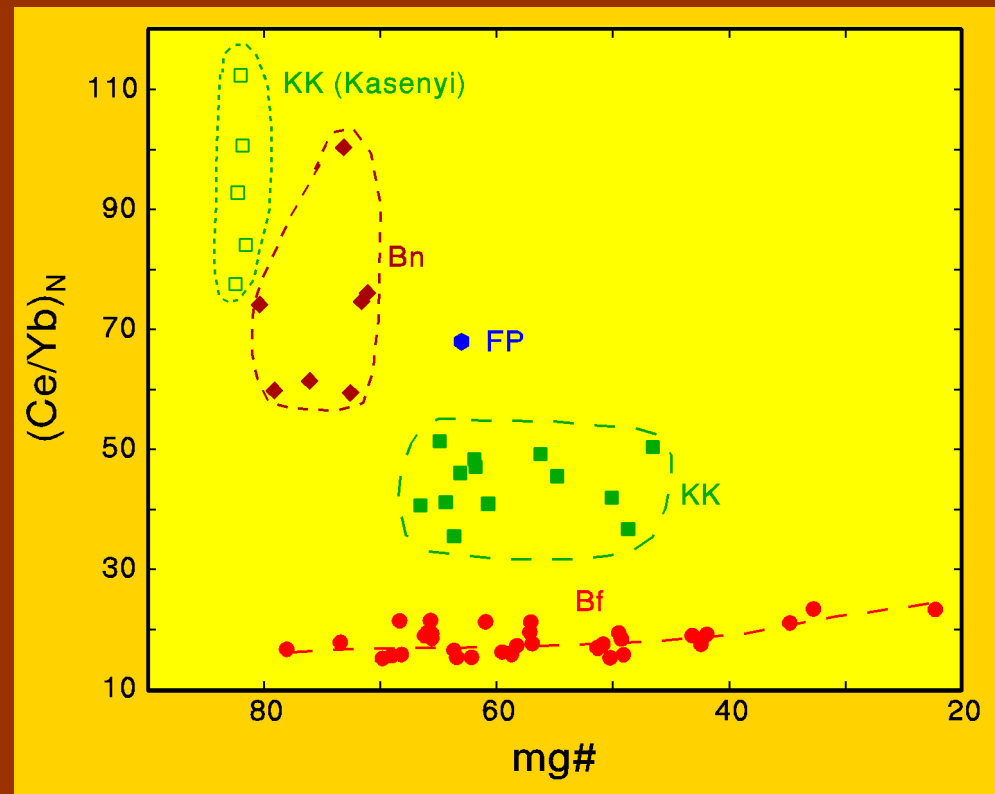


Zr/Hf and Nb/Ta ratios are close to mantle values with the exception of Fort Portal and some of the K-rich samples which show relative Zr and/or Nb enrichment.

OIB-normalized spider diagrams for the various volcanic fields. Note the differences in the slopes for the various spider diagrams.



$(\text{Ce}/\text{Yb})_N$ ratios as a function of mg#. The high-K samples have the highest $(\text{Ce}/\text{Yb})_N$ ratios suggesting smaller degrees of partial melting. Bufumbira samples show a regular increase in $(\text{Ce}/\text{Yb})_N$ with decreasing mg#, a trend typical of fractional crystallization.



Representative REE patterns. Note crossing patterns for Katwe-Kikorongo, Bunyaraguru and Bufumbira which suggest that the melts were derived by variable degrees of melting of a garnet-bearing source.

Conclusions

- The mantle under the western branch of the East African rift system is chemically heterogeneous.
- This mantle has been subjected to variable degrees of metasomatism as represented by the K_2O/Na_2O ratios for volcanics from the various fields.
- Chemical mantle domains can be identified which gave rise to each of the volcanic groups.
- REE data suggest that garnet was important in the source area, and that magma variations are in part due to variable degrees of partial melting.