

# Notes on The Laplace Transform and the Ozone Layer

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## 1 The Laplace Transform and the Ozone Layer

In farfield propagation problems, we often find the measured data to be related to the desired object function by a Fourier transformation. The image reconstruction problem then becomes one of estimating a function from finitely many noisy values of its Fourier transform. In this note we consider an inverse problem involving the Laplace transform. The example is taken from Twomey's book [1].

## 2 The Laplace Transform

The Laplace transform of the function  $f(x)$  defined for  $0 \leq x < +\infty$  is the function

$$\mathcal{F}(s) = \int_0^{+\infty} f(x)e^{-sx} dx.$$

## 3 Scattering of Ultraviolet Radiation

The sun emits ultraviolet (UV) radiation that enters the Earth's atmosphere at an angle  $\theta_0$  that depends on the sun's position, and with intensity  $I(0)$ . Let the  $x$ -axis be vertical, with  $x = 0$  at the top of the atmosphere and  $x$  increasing as we move down to the Earth's surface, at  $x = X$ . The intensity at  $x$  is given by

$$I(x) = I(0)e^{-kx/\cos\theta_0}.$$

Within the ozone layer, the amount of UV radiation scattered in the direction  $\theta$  is given by

$$S(\theta, \theta_0)I(0)e^{-kx/\cos\theta_0} \Delta p,$$

where  $S(\theta, \theta_0)$  is a known parameter, and  $\Delta p$  is the change in the pressure of the ozone within the infinitesimal layer  $[x, x + \Delta x]$ , and so is proportional to the concentration of ozone within that layer.

## 4 Measuring the Scattered Intensity

The radiation scattered at the angle  $\theta$  then travels to the ground, a distance of  $X - x$ , weakened along the way, and reaches the ground with intensity

$$S(\theta, \theta_0)I(0)e^{-kx/\cos\theta_0}e^{-k(X-x)/\cos\theta}\Delta p.$$

The total scattered intensity at angle  $\theta$  is then a superposition of the intensities due to scattering at each of the thin layers, and is then

$$S(\theta, \theta_0)I(0)e^{-kX/\cos\theta_0}\int_0^Xe^{-x\beta}dp,$$

where

$$\beta = k\left[\frac{1}{\cos\theta_0} - \frac{1}{\cos\theta}\right].$$

This superposition of intensity can then be written as

$$S(\theta, \theta_0)I(0)e^{-kX/\cos\theta_0}\int_0^Xe^{-x\beta}p'(x)dx.$$

## 5 The Laplace Transform Data

Using integration by parts, we get

$$\int_0^Xe^{-x\beta}p'(x)dx = p(X)e^{-\beta X} - p(0) + \beta\int_0^Xe^{-\beta x}p(x)dx.$$

Since  $p(0) = 0$  and  $p(X)$  can be measured, our data is then the Laplace transform value

$$\int_0^{+\infty}e^{-\beta x}p(x)dx;$$

note that we can replace the upper limit  $X$  with  $+\infty$  if we extend  $p(x)$  as zero beyond  $x = X$ .

The variable  $\beta$  depends on the two angles  $\theta$  and  $\theta_0$ . We can alter  $\theta$  as we measure and  $\theta_0$  changes as the sun moves relative to the earth. In this way we get values of the Laplace transform of  $p(x)$  for various values of  $\beta$ . The problem then is to recover  $p(x)$  from these values. Because the Laplace transform involves a smoothing of the function  $p(x)$ , recovering  $p(x)$  from its Laplace transform is more ill-conditioned than is the Fourier transform inversion problem.

## References

- [1] Twomey, S. (1996) *Introduction to the Mathematics of Inversion in Remote Sensing and Indirect Measurement*. New York: Dover Publ.