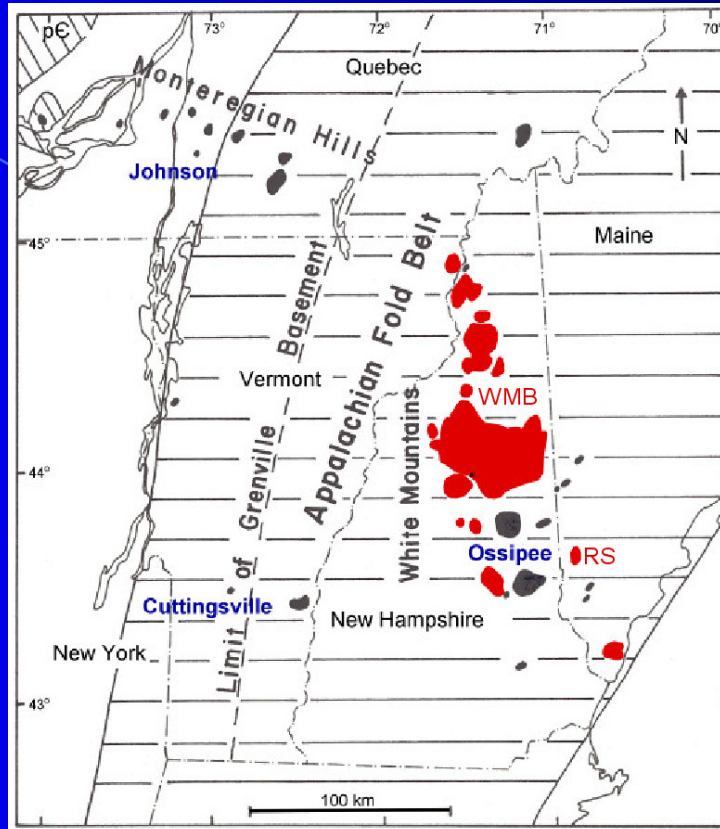
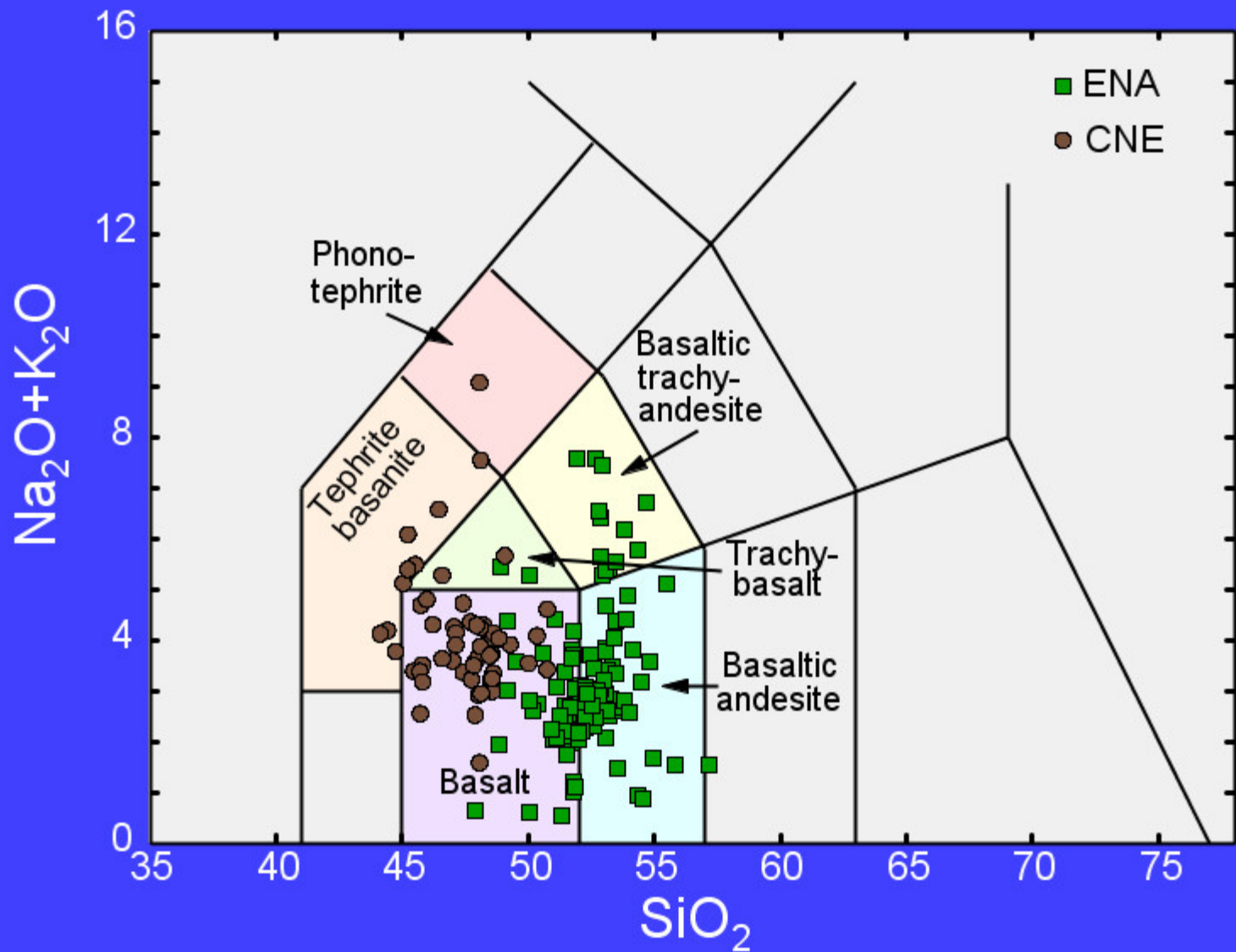


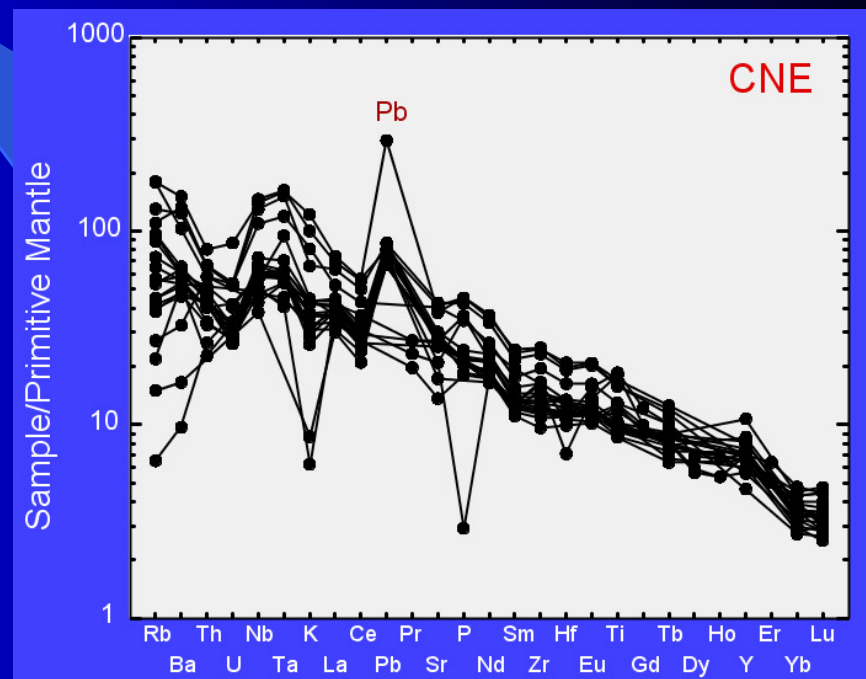
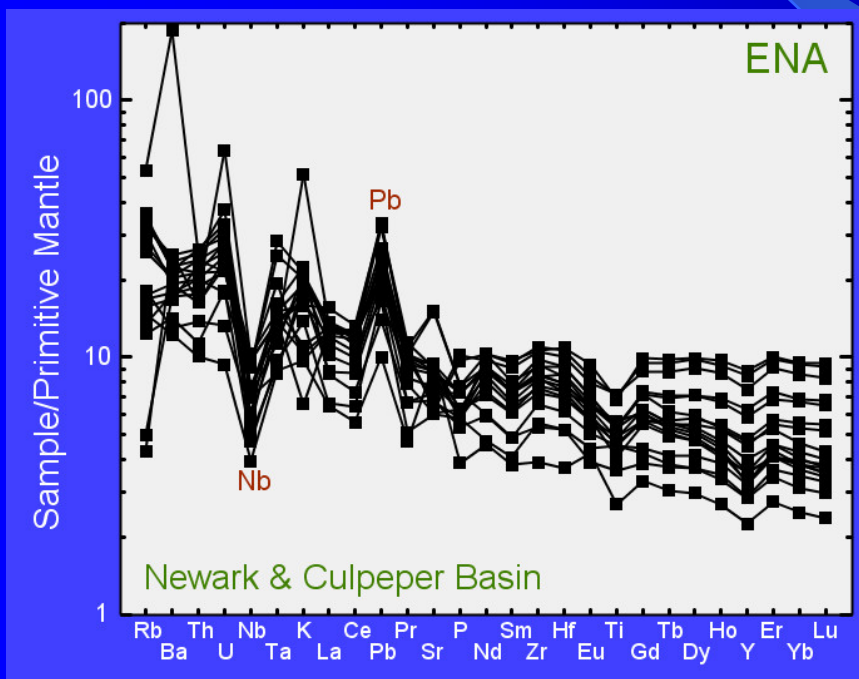
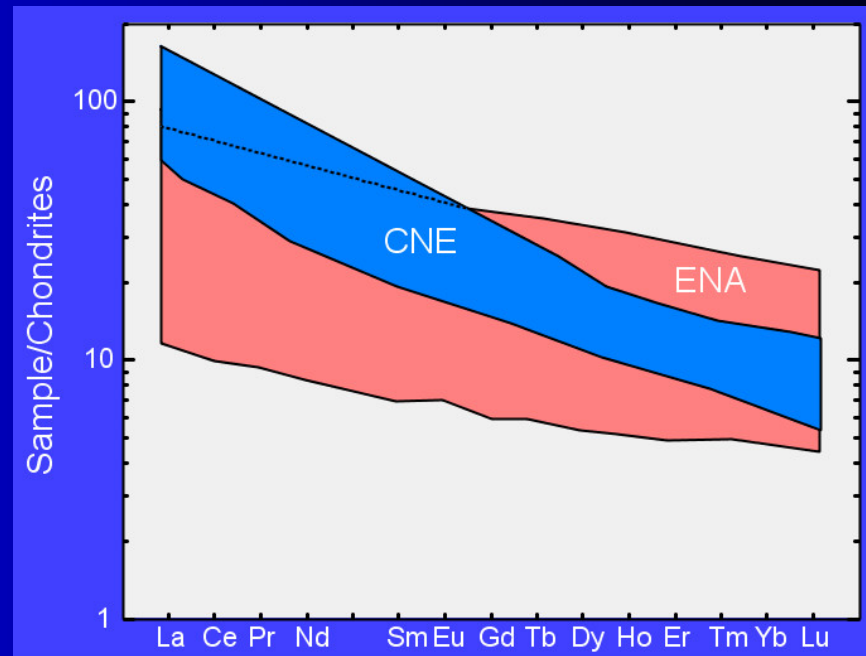
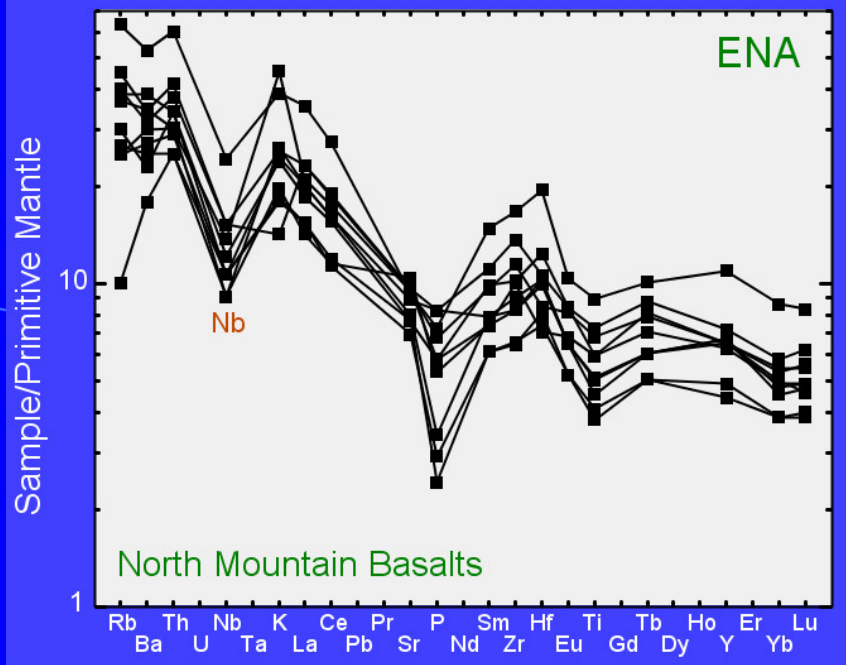
Post CAMP Magmatism: The White Mountain and Montereian Hills Igneous Provinces, Eastern North America



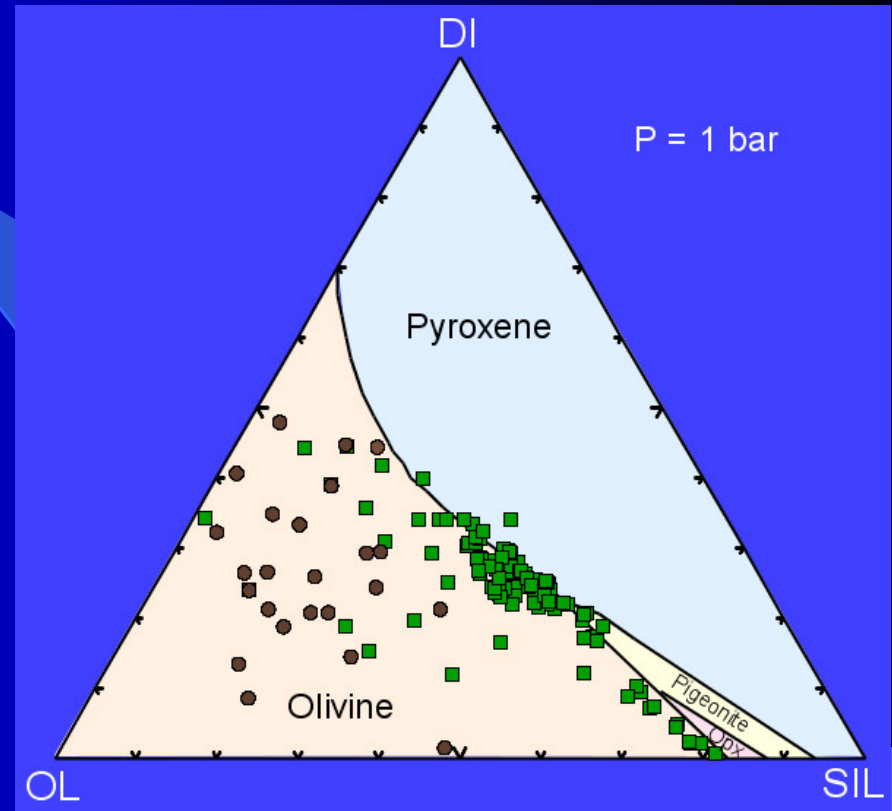
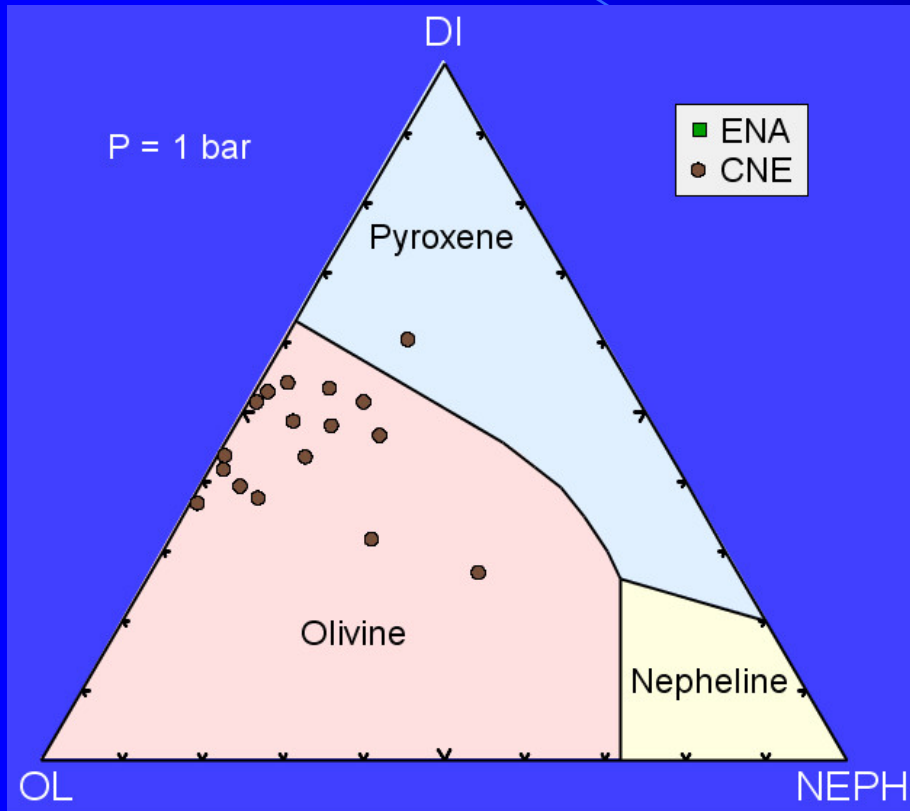
Nelson Eby
Department of Environmental, Earth & Atmospheric Sciences
University of Massachusetts, Lowell

Comparison of the Coastal New England (CNE) and Eastern North America (ENA) suites

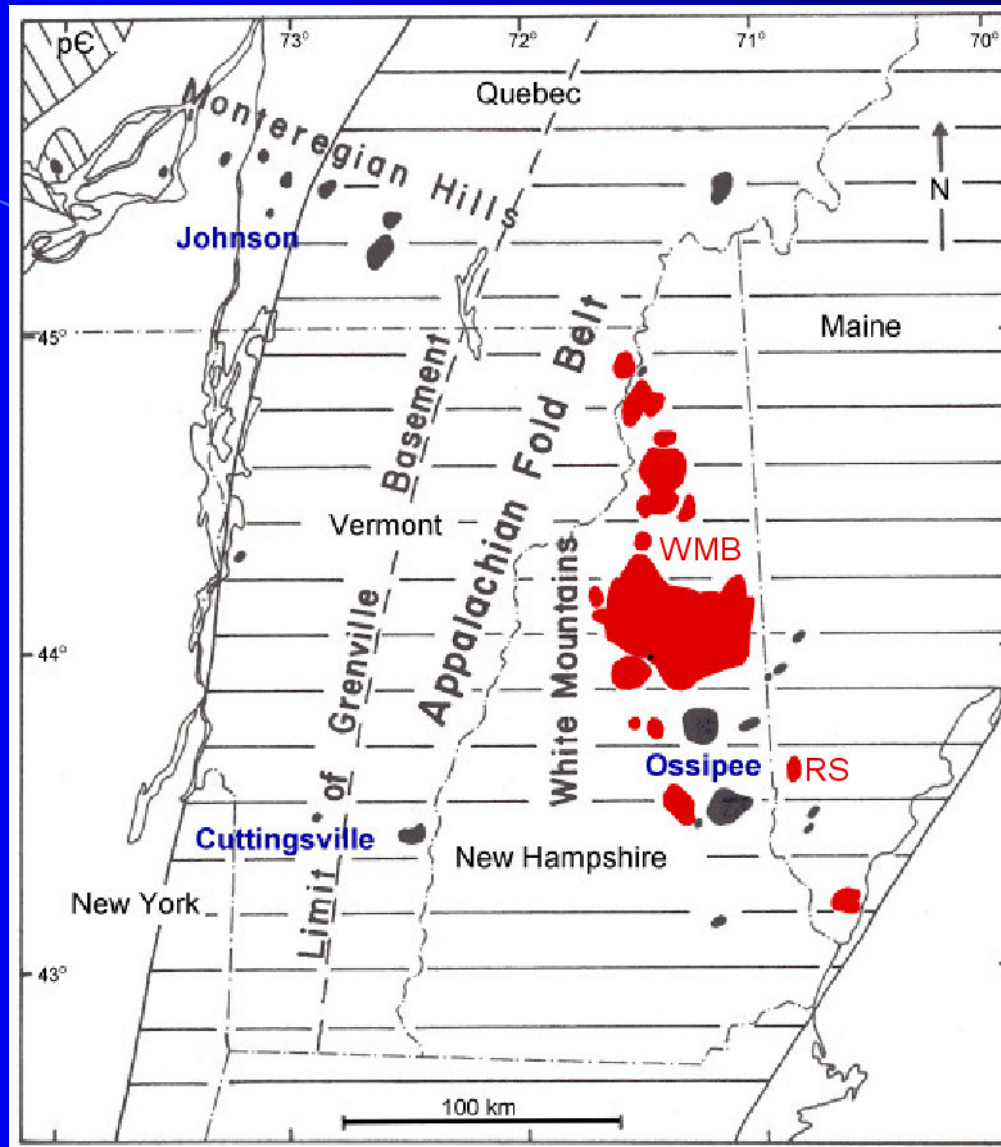


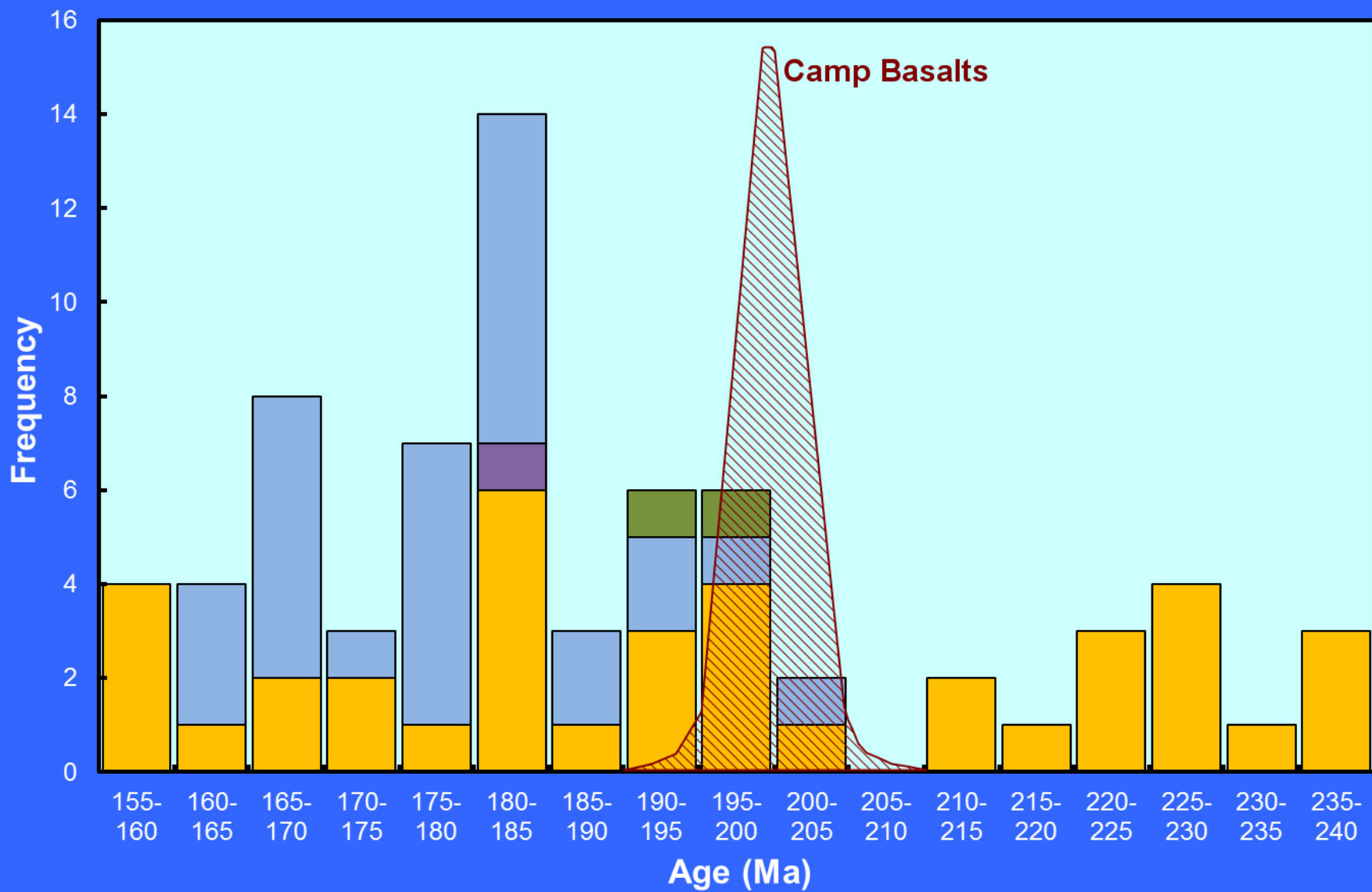


CNE and ENA basalt compositions projected into 1 bar phase diagrams. The ENA basalts plot on the olivine-pyroxene cotectic. With one exception, on both phase diagrams the CNE basalts plot in the olivine field. Increasing pressure shifts the olivine-pyroxene cotectic towards the olivine corner. The presence of both olivine and pyroxene as liquidus phases would indicate high pressure fractionation.

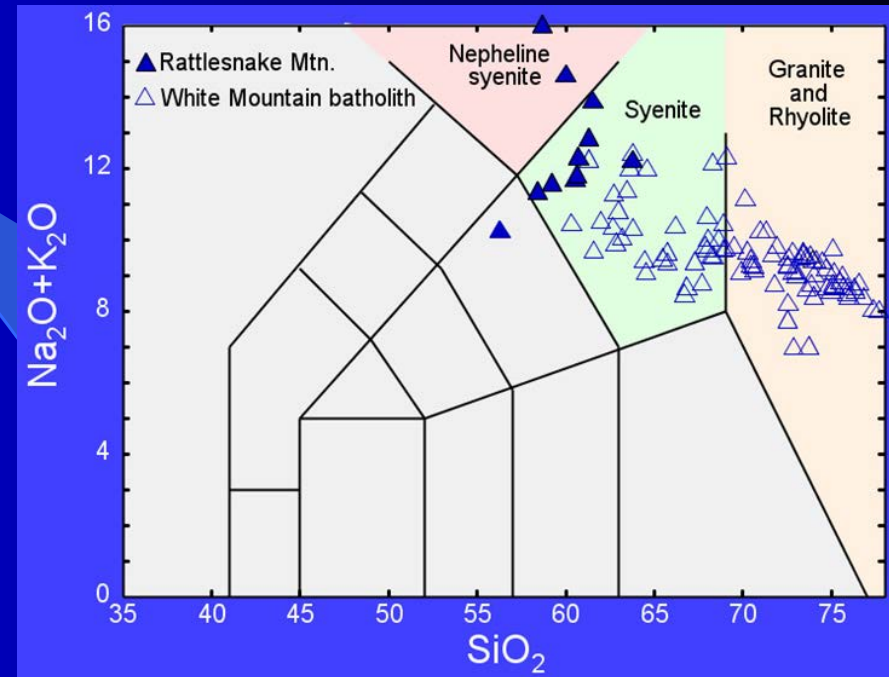
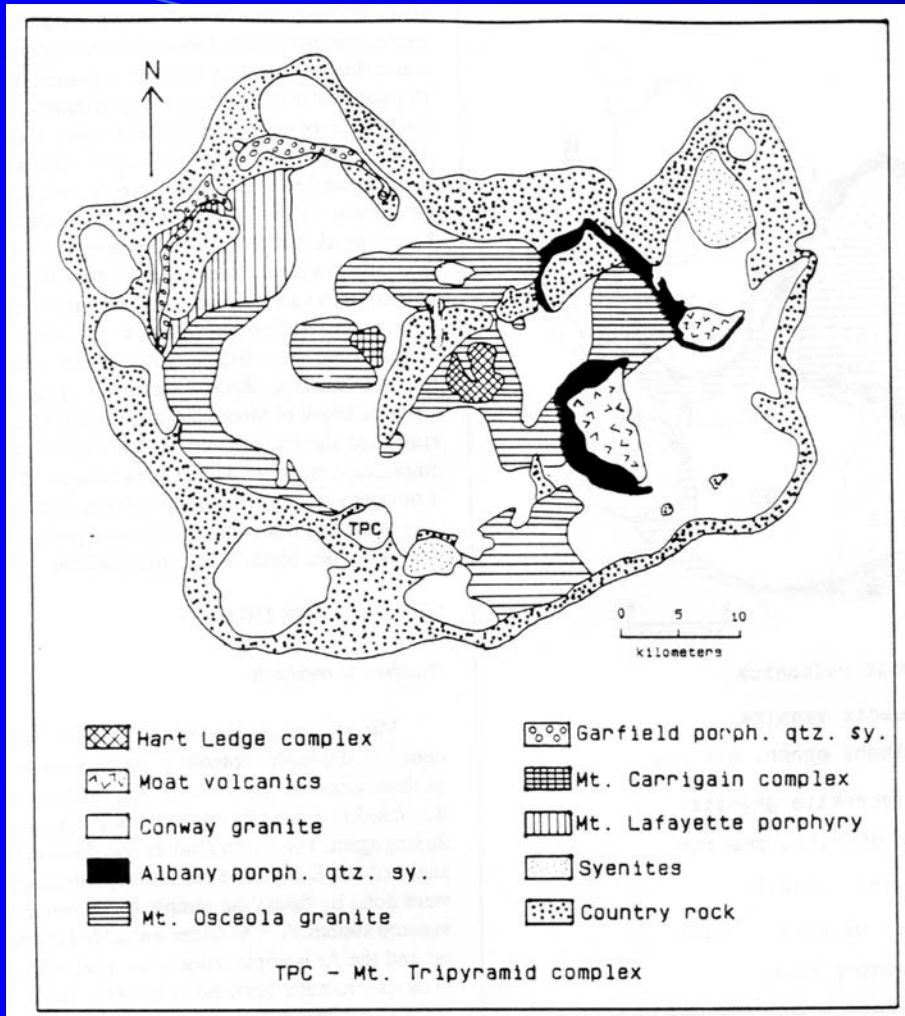


Comparison of the Coastal New England (CNE) and Older White Mountain (OWM) suites

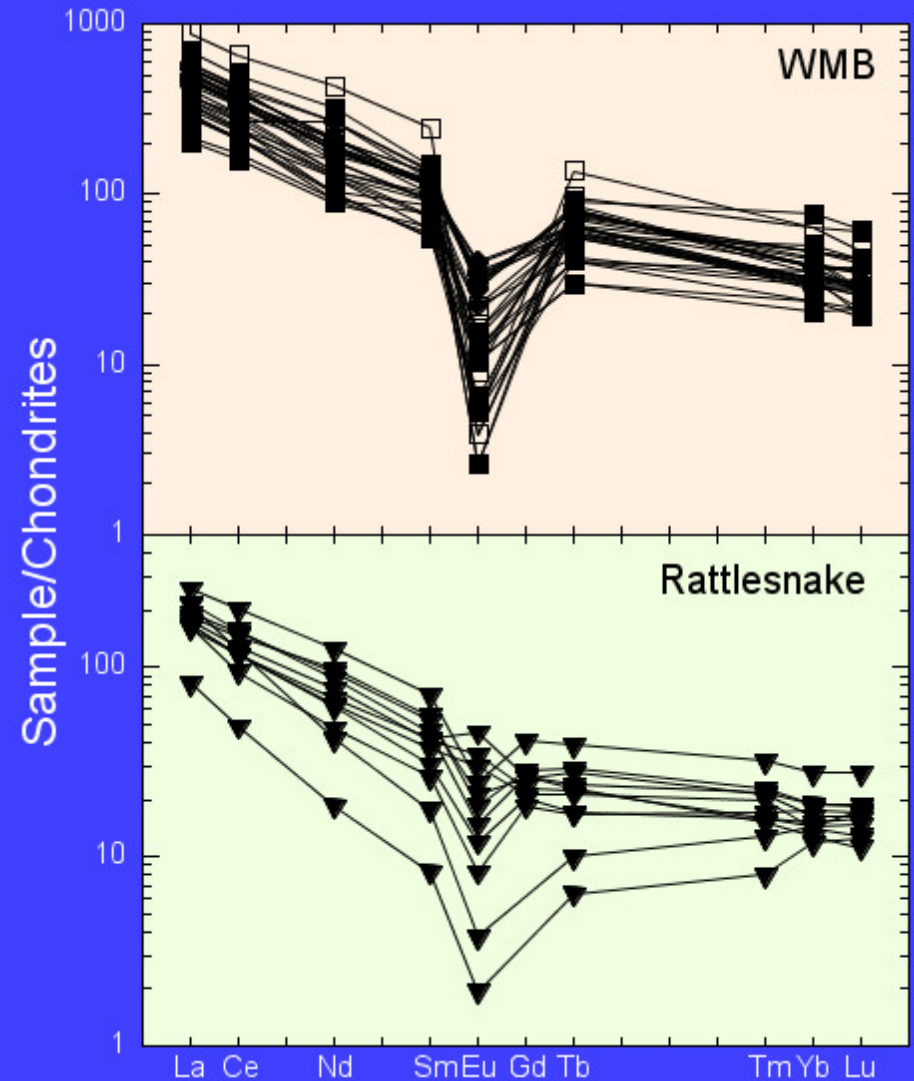
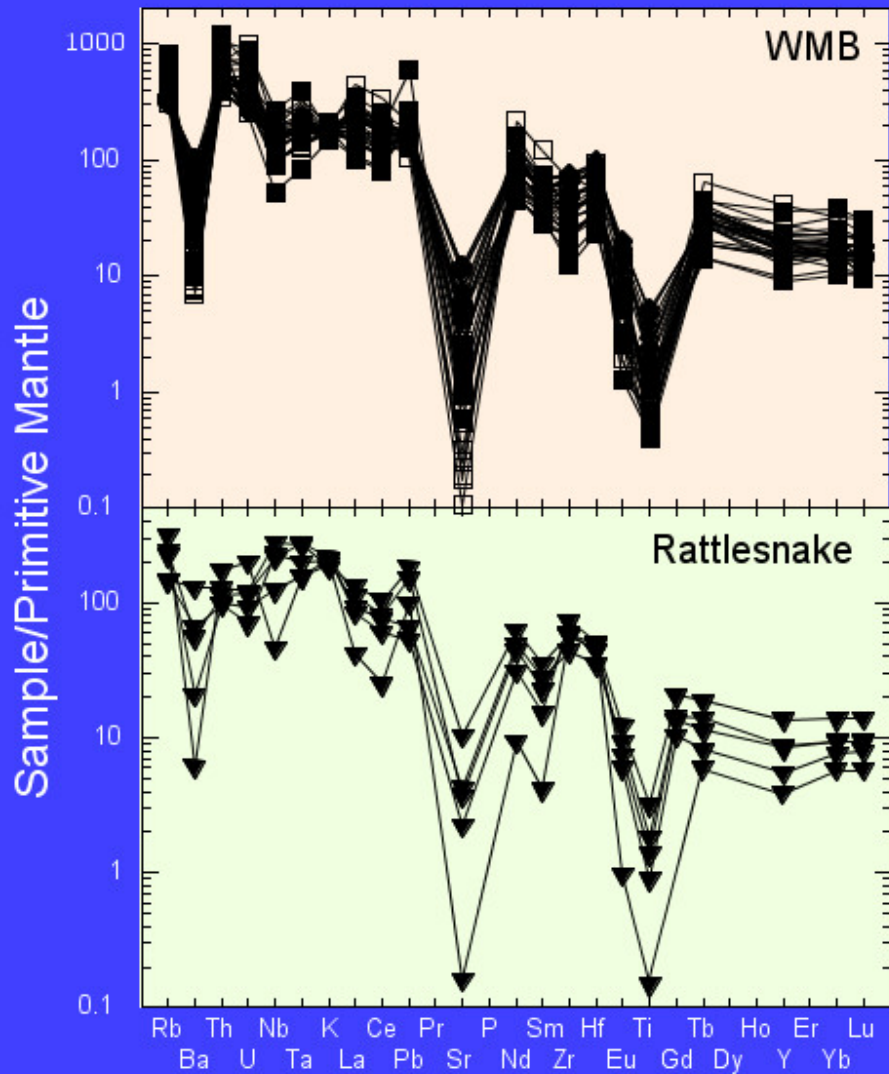




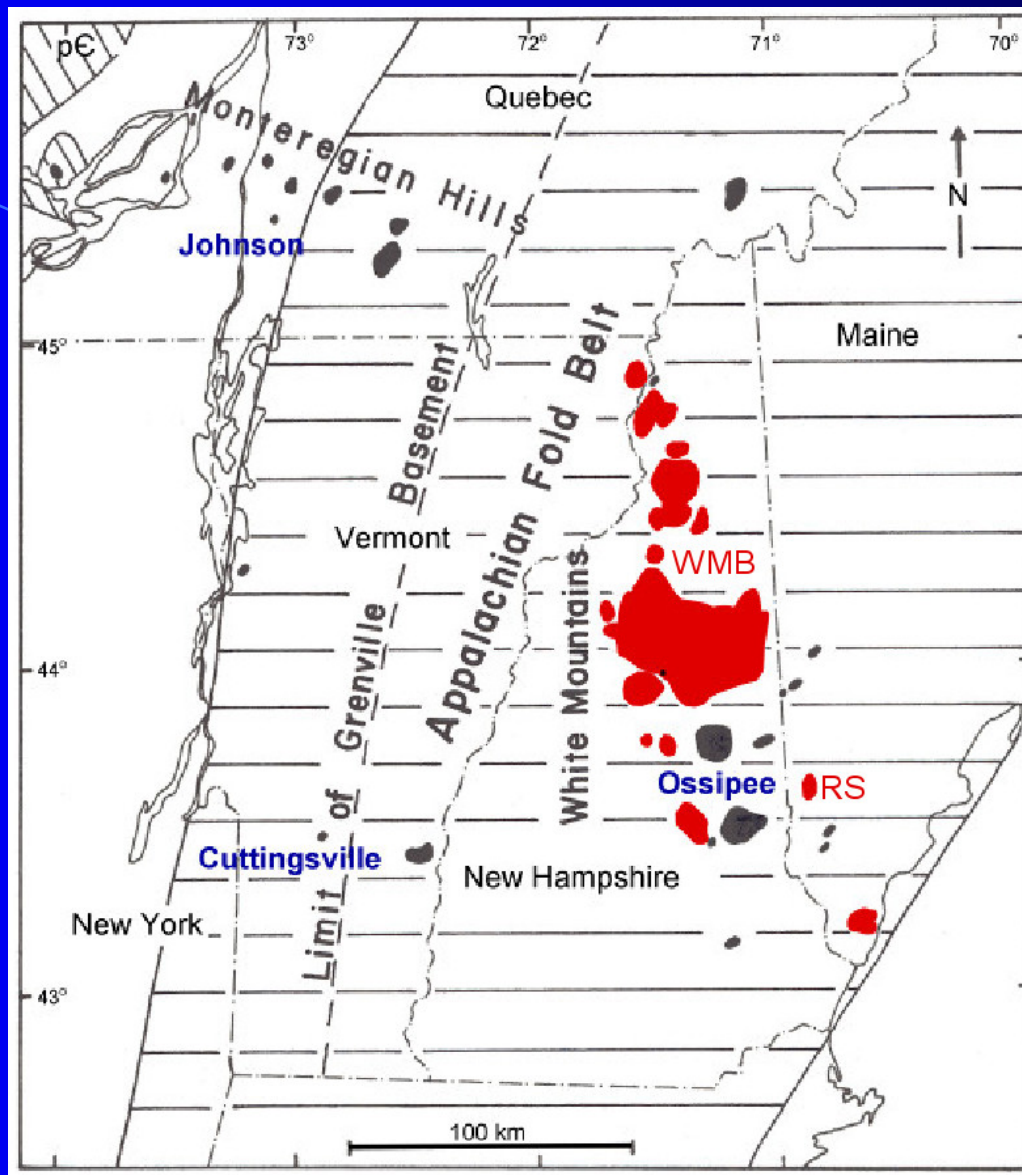
The White Mountain batholith is comprised of syenites, granites, and rhyolites. The Rattlesnake pluton consists of nepheline syenite and syenite. Mafic rocks are absent. In general, mafic rocks are not significant components of the Older White Mountain igneous sequences.

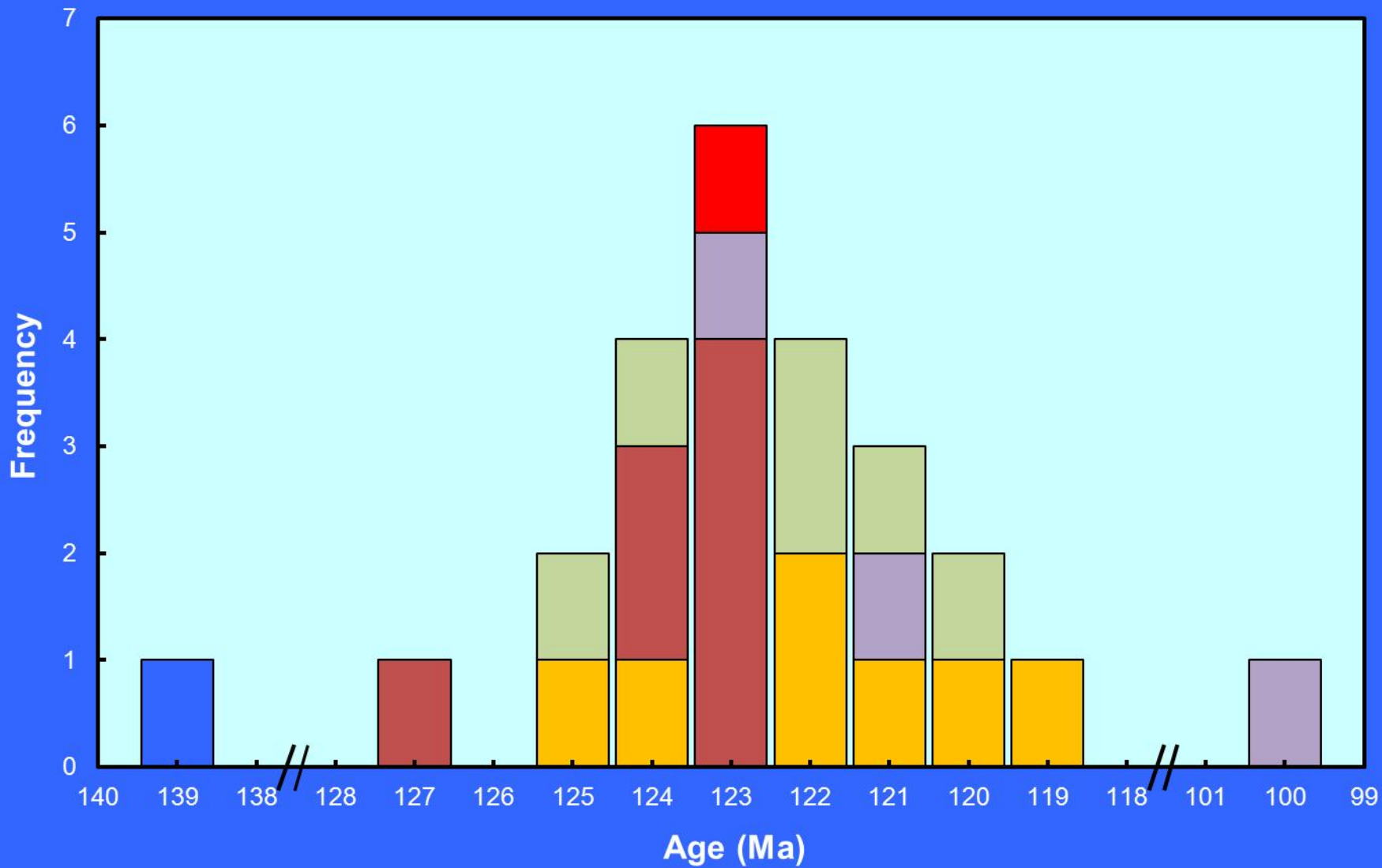


Significant depletions in Ba, Sr, and Eu indicate significant feldspar fractionation. Of note is the small positive Pb anomaly and the small negative Nb anomaly.



Monteregian Hills and Younger White Mountains – a Tale of Three Plutons – Ossipee, Cuttingsville, and Ossipee.



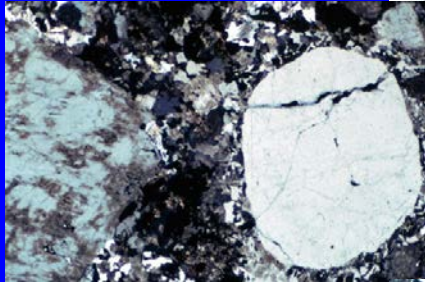
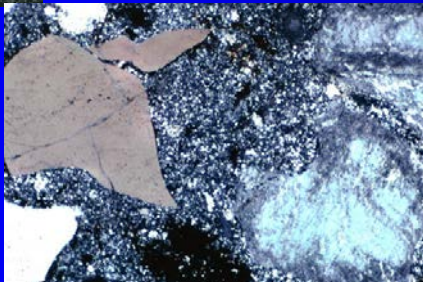


Ossipee ring complex



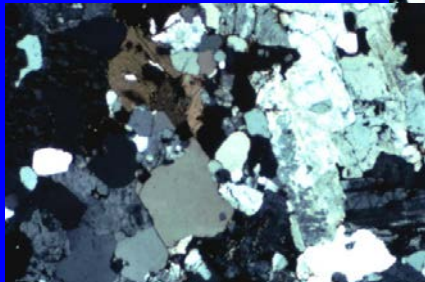
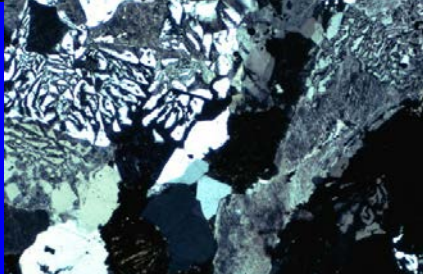
Basalt

Rhyolite

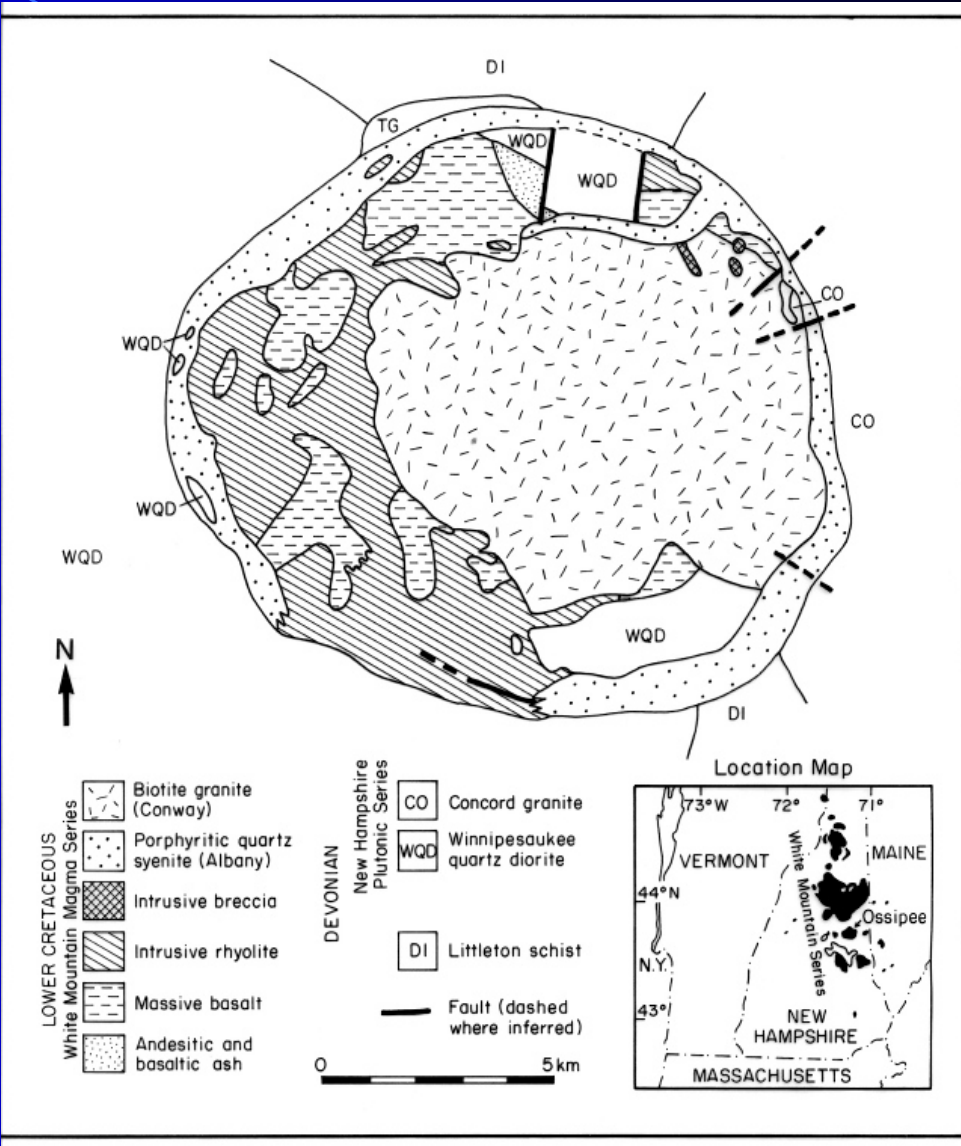


Porphyritic quartz syenite

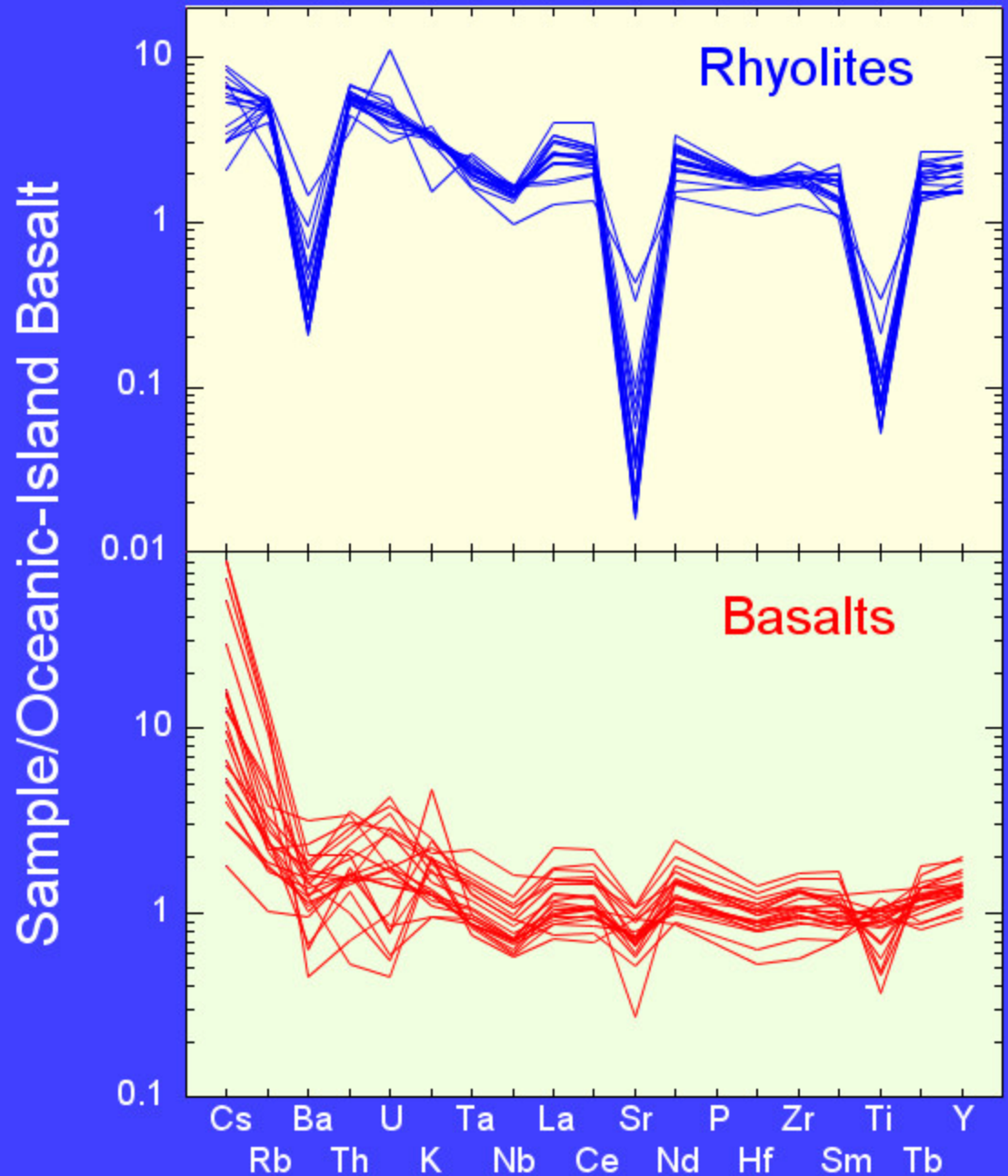
Microgranite



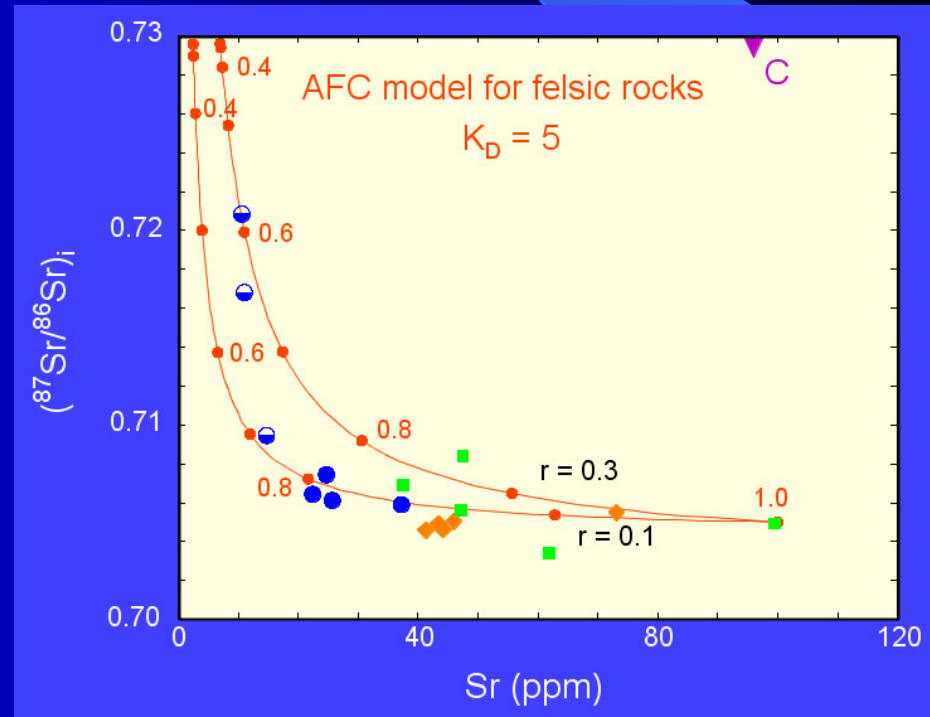
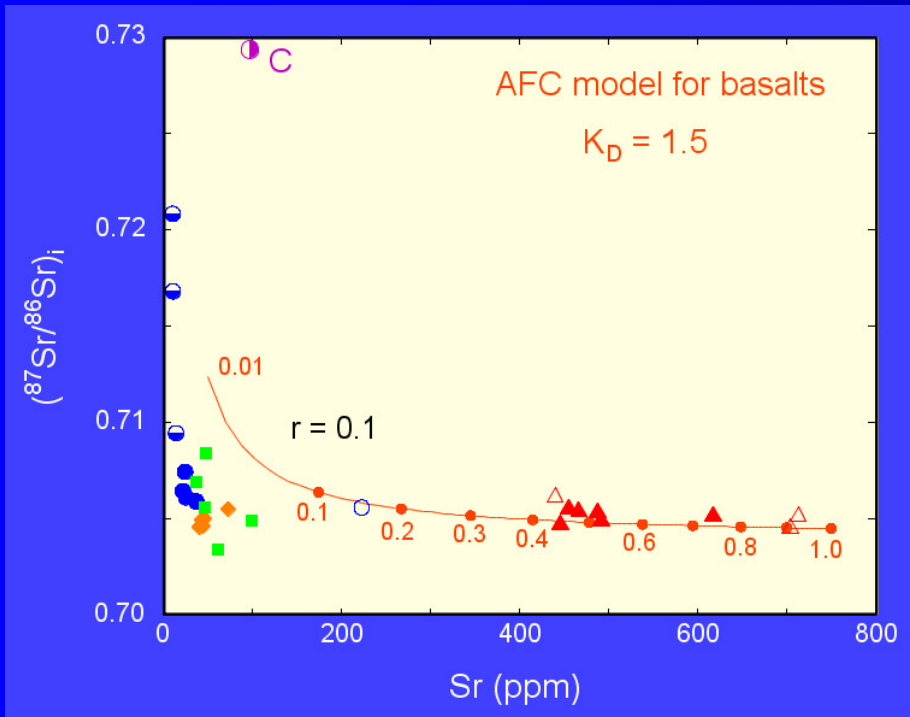
Biotite granite



OIB normalized spider diagrams for Ossipee rhyolites and basalts. Note the similarity of both lithologies to OIB. Variations can be explained by the fractionation of alkali feldspar and opaque oxide minerals. Cs enrichment in basalts is due to late-stage hydrothermal alteration as evidence by the partial replacement of plagioclase by epidote.



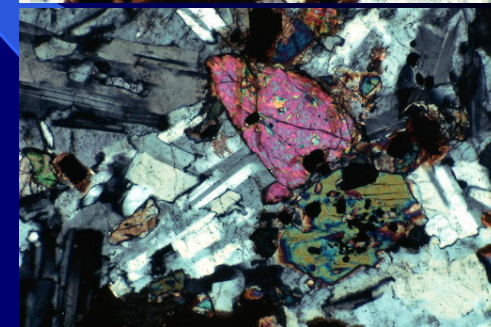
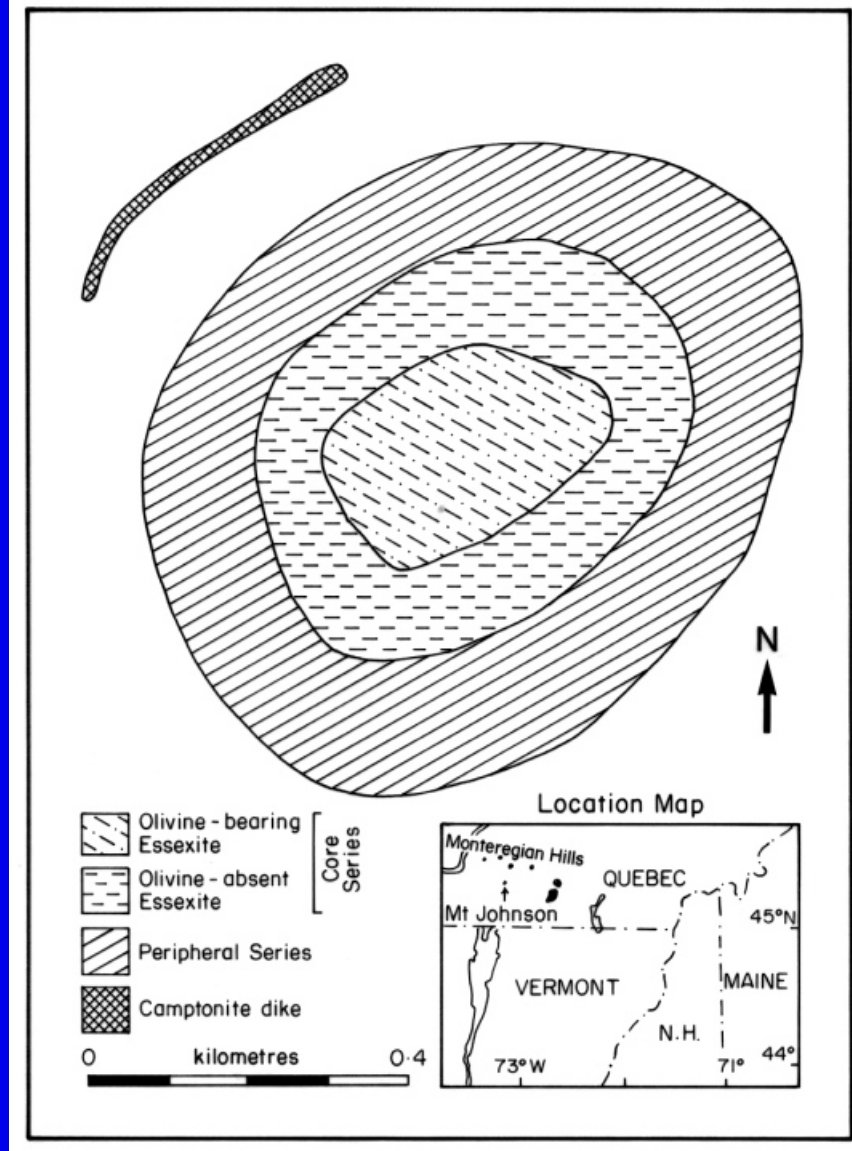
AFC models for basalts and felsic rocks. The isotopic variations require only minor contamination of the melts by country rock.



Mont Johnson

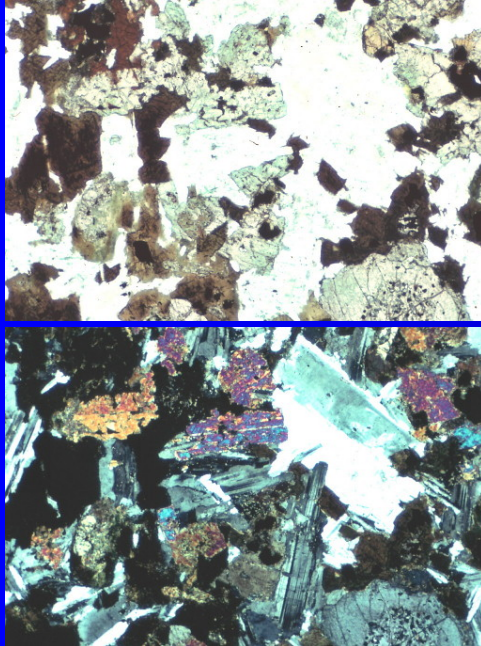


Core series
essexite from
center of pluton.

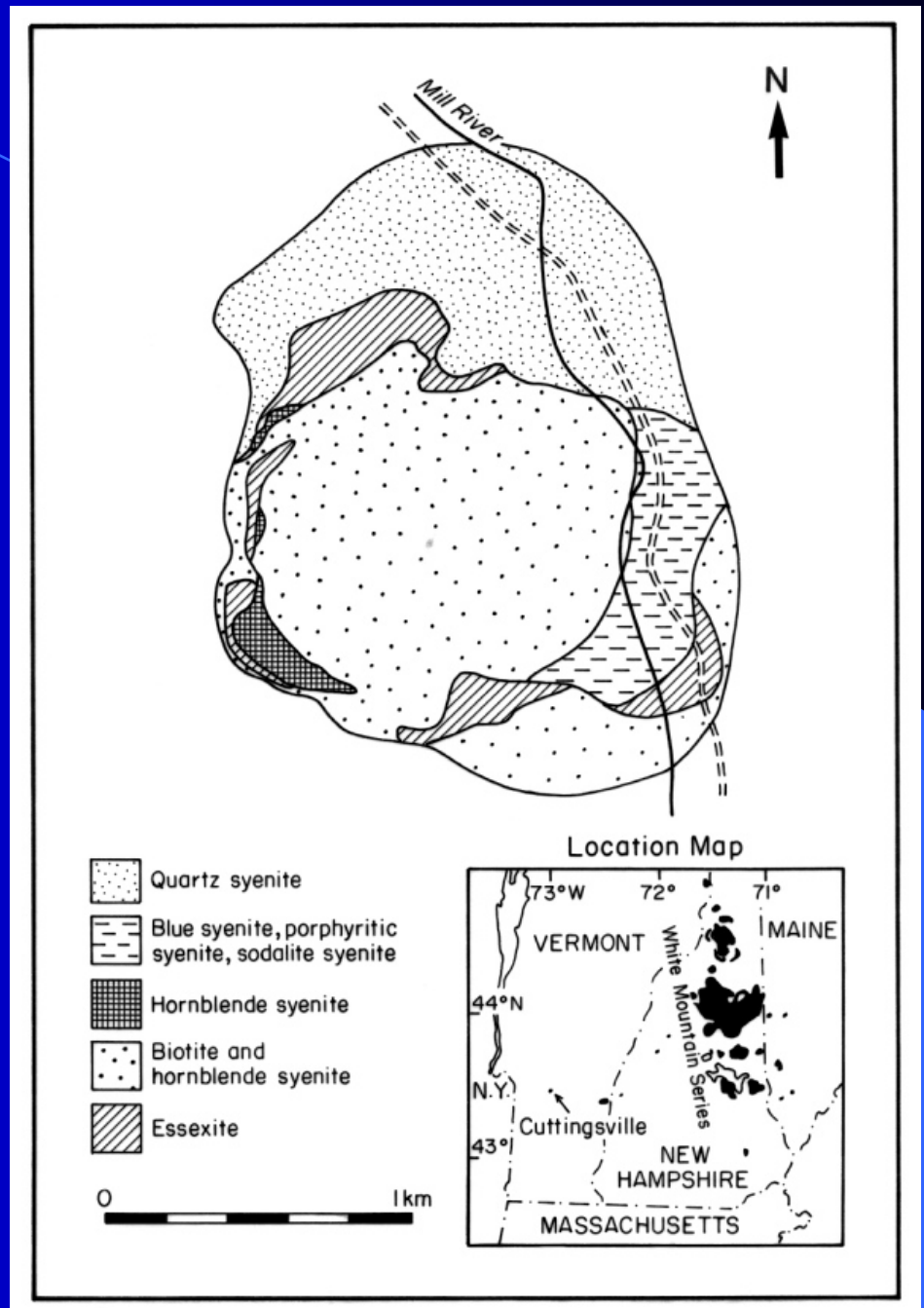


Core series
essexite. Width of
field of view,
2mm.

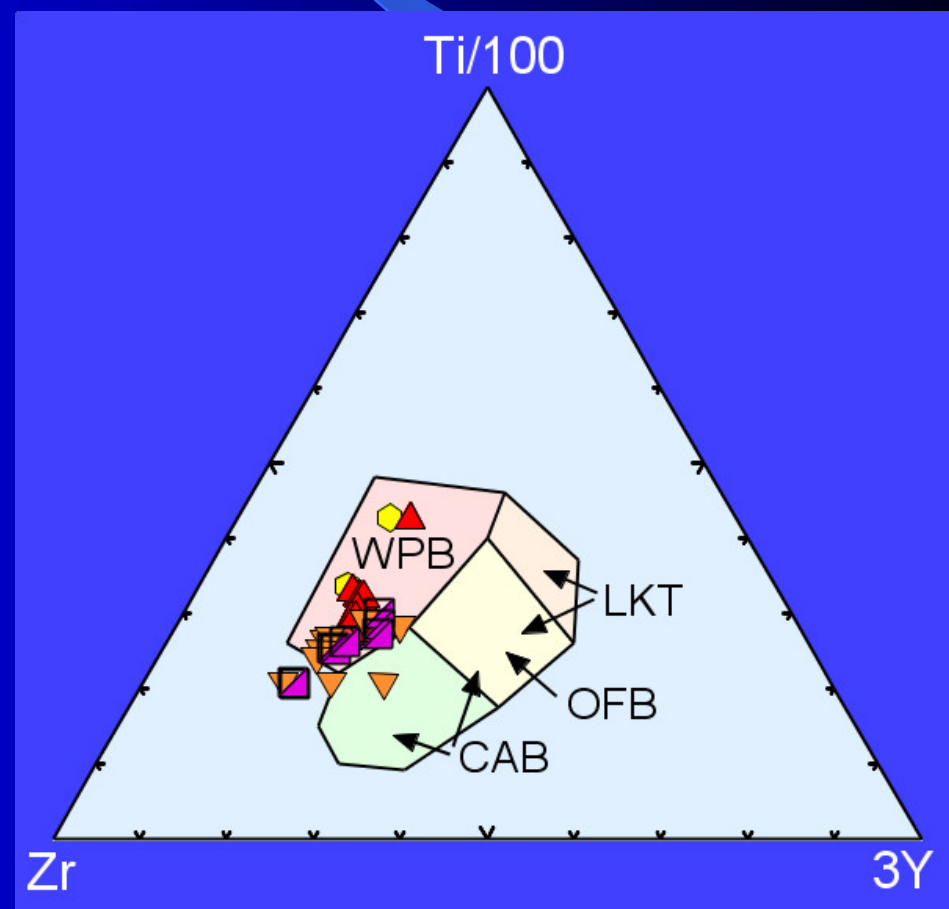
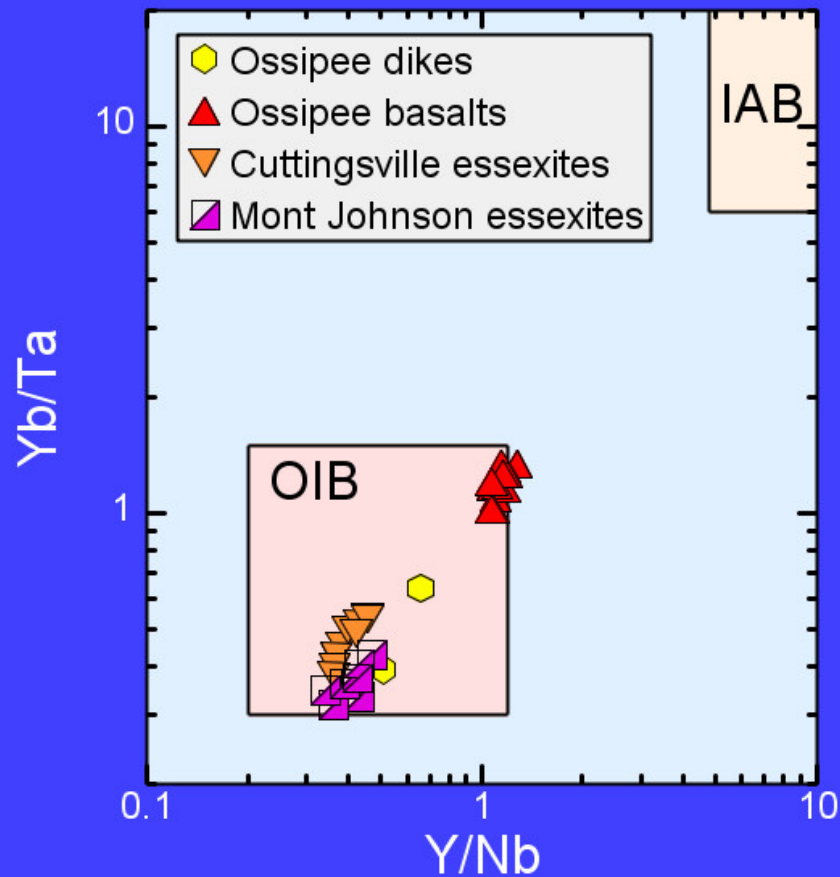
Cuttingsville



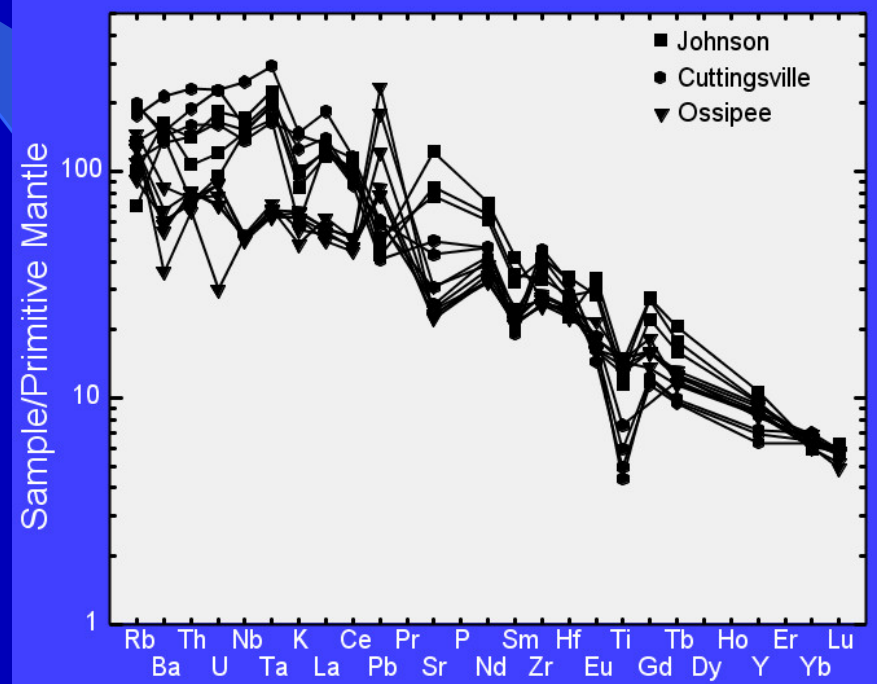
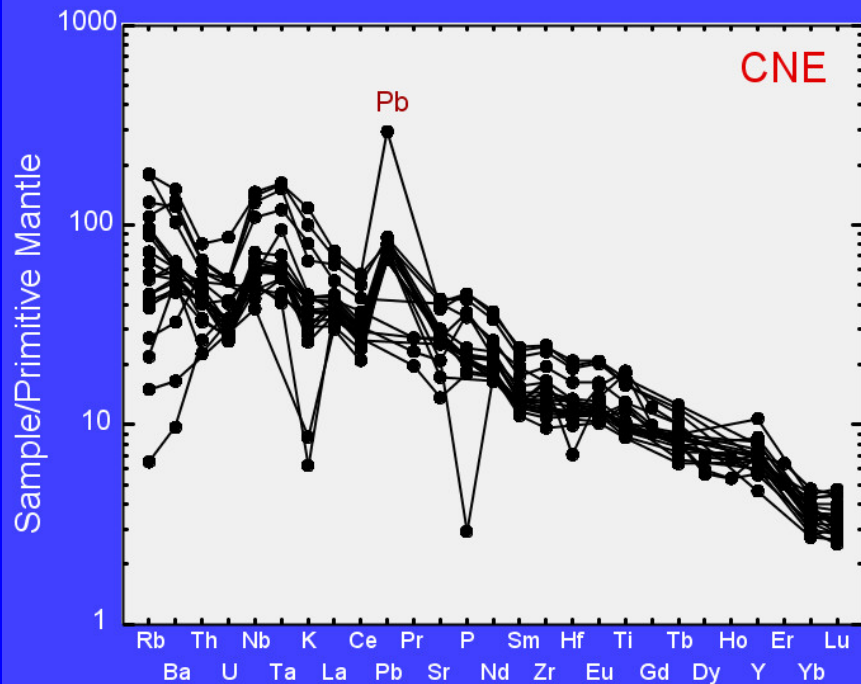
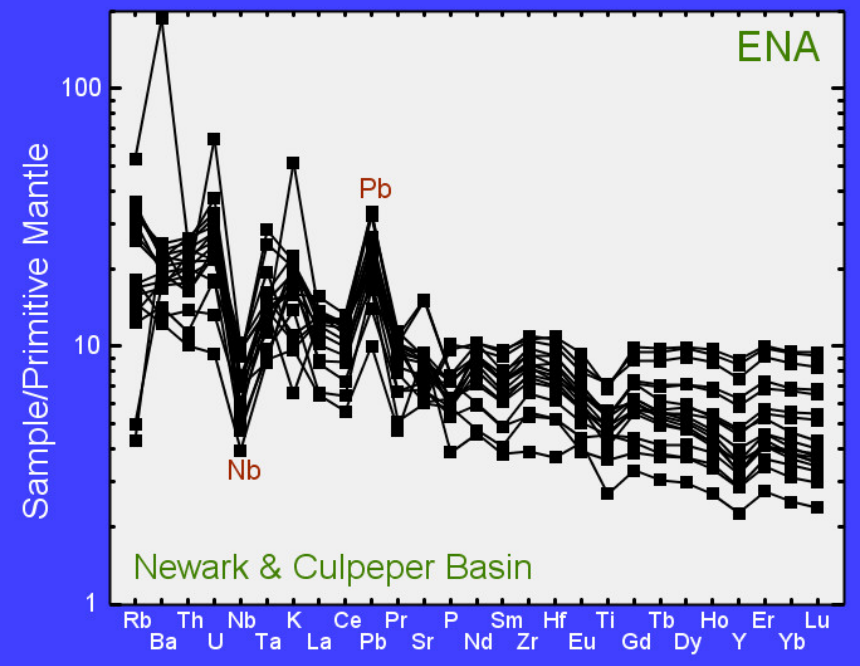
Essexite



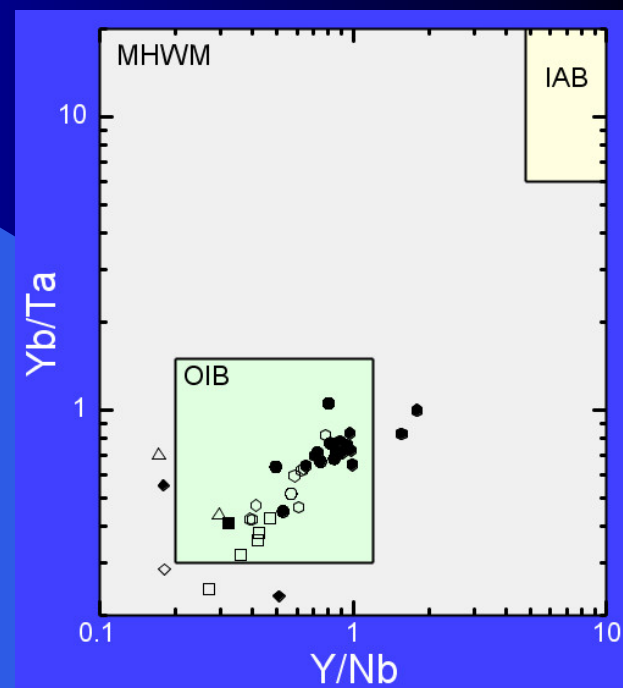
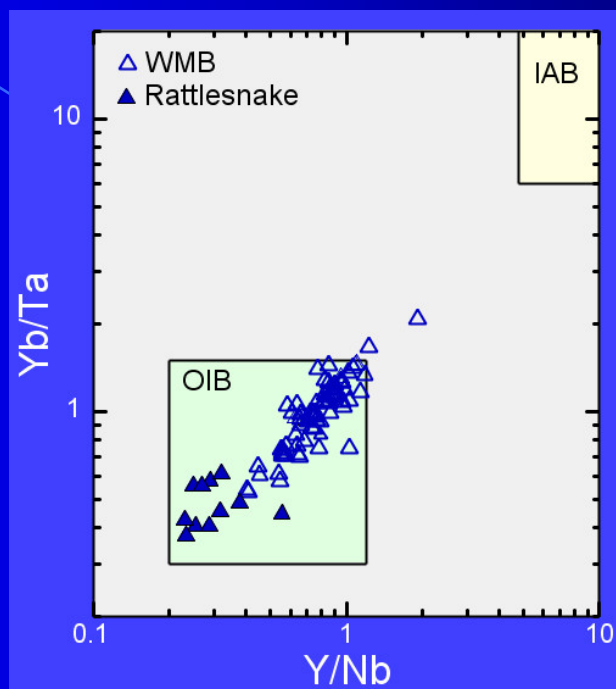
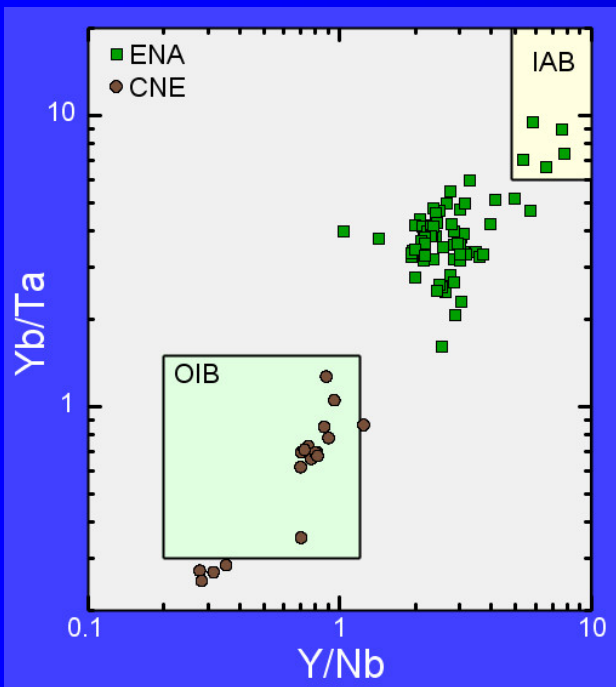
The mafic rocks plot in the OIB and WPB fields on various discrimination diagrams. In the Y/Nb vs Yb/Ta diagram the Ossipee basalts plot towards the IAB field (but still within the OIB field), an indication of minor crustal contamination.



Comparison of ENA, CNE and MHWM elemental chemistry. Note the Nb depletion in ENA. CNE and MHWM patterns show Nb depletions and enrichments. Prominent positive Pb anomaly in ENA and CNE, but absent in MHWM except for Ossipee which was emplaced into the folded Appalachians. The Pb enrichment *may* represent crustal interaction.



Y/Nb versus Yb/Ta. CNE, OWM and MHWM all fall in the OIB field. The ENA samples plot towards and in the IAB field.



Melting models for various mantle sources. Note that the MHWM mafic rocks fall along the Garnet Peridotite (GP) curve and are apparently related by variable degrees of melting of the same source. The ENA basalts plot in the general area of melts formed in equilibrium with a Depleted Spinel (SD) lherzolite mantle. The CNE dikes fall between the curves perhaps suggesting a mixed garnet-spinel source.

